

# Geography of the world's oceans and major current systems

Lecture 2

WHY is the GEOMORPHOLOGY OF THE  
OCEAN FLOOR important?

(in the context of Oceanography)

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(in the context of Oceanography)

- Ocean circulation, tides, and mixing on regional and basin scales are heavily controlled by the topography of the ocean.
- The nature of the earth, its origin, and its characteristics have a profound effect on the properties and the composition of the biota that are contained in the ocean.
- The structure and distribution of sediments can be understood based on the geomorphology of the ocean floor. These sediments are important because they tell us about the geochemistry of the ocean floor. Also they can be used to reconstruct ocean circulation of the past and improve our understanding of the climate system.

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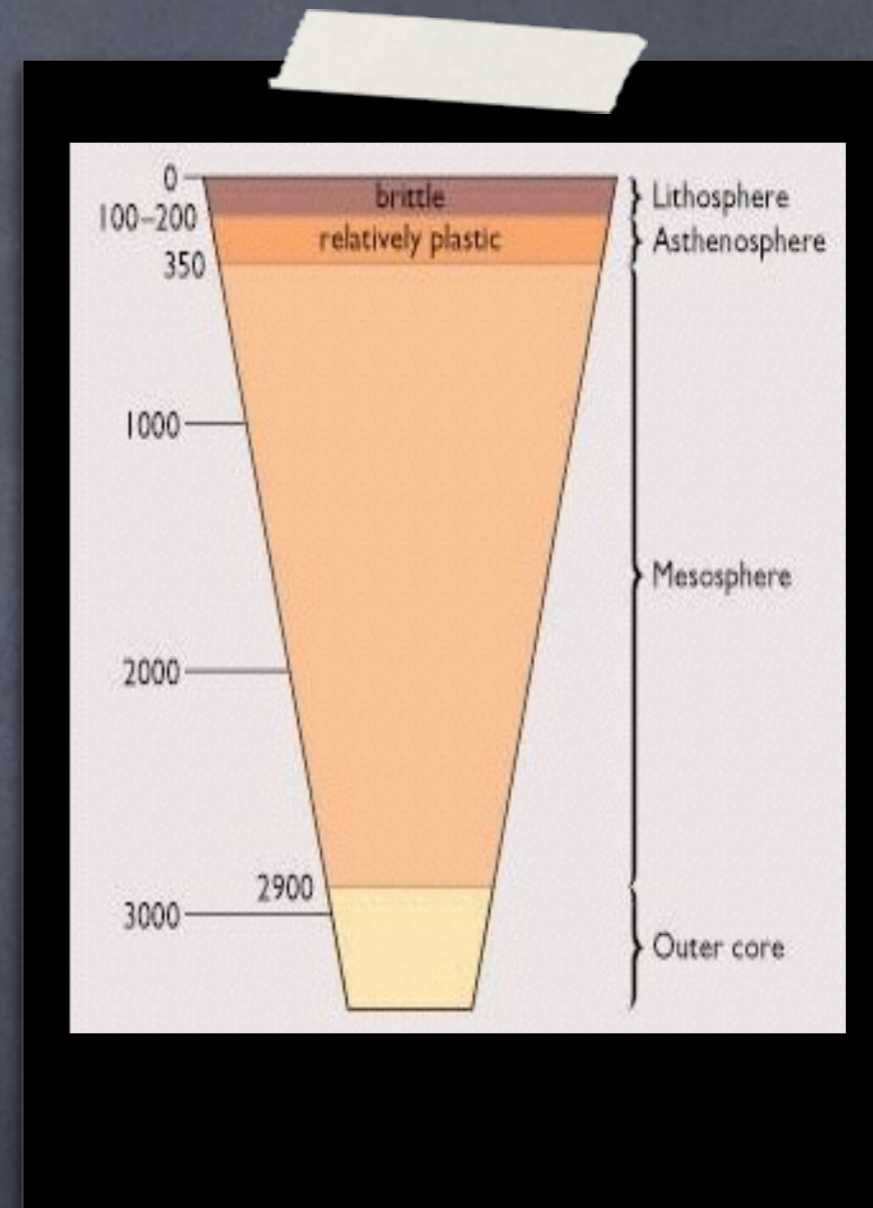
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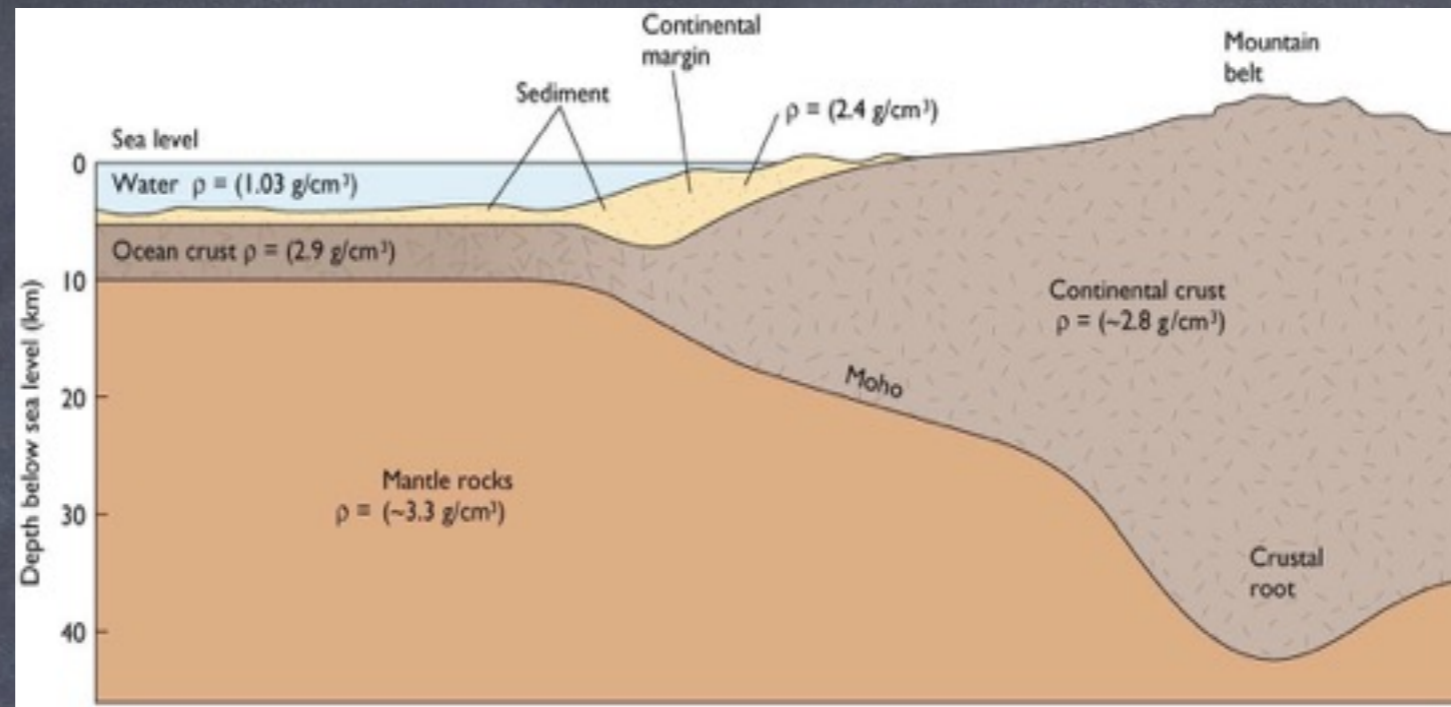
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Divisions of the Earth based upon physical state are:

- **Lithosphere** or crust (hard rock: pressure effect dominates)
- **Asthenosphere** (mantle, <1% of rock melted, similar to tar or asphalt: temperature effect significant)
- **Mesosphere** (more rigid part of deeper Mantle: pressure effect dominates)
- **Outer and inner core** (molten and solid, both made up of iron and nickel alloy)



# Geologic Differences between Continents and Ocean Basins

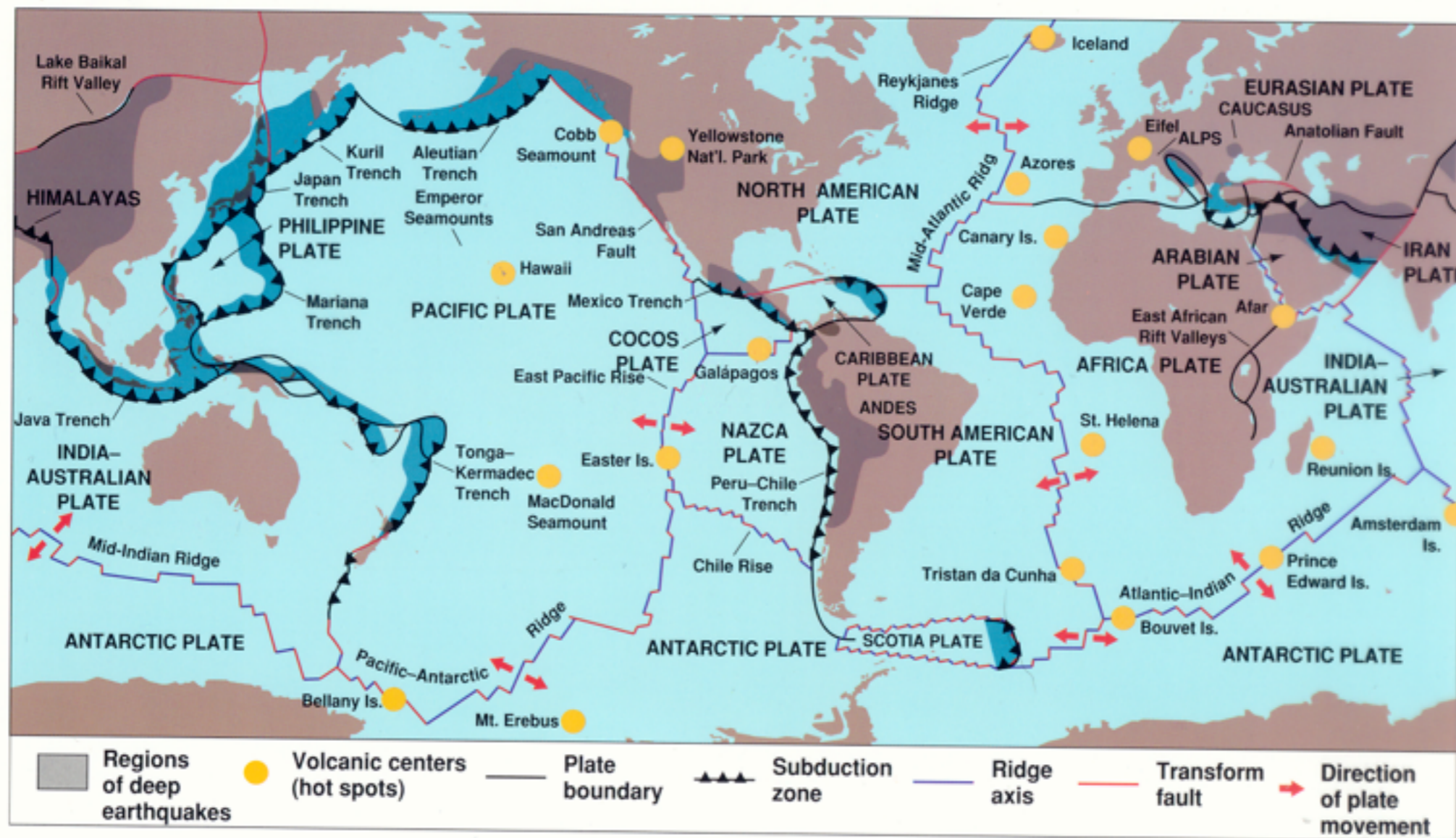


- **Continental crust** is mainly composed of granite, a light colored, lower density ( $2.8 \text{ gm/cm}^3$ ) igneous rock rich in aluminum, silicon and oxygen.
- **Oceanic crust** is composed of basalt, a dark colored, higher density ( $2.9 \text{ gm/cm}^3$ ) volcanic rock rich in silicon, oxygen and magnesium.
- **Oceanic crust is thin and dense. Continental crust is thick and light.**

**Isostasy** is a term used in Geology to refer to the state of gravitational equilibrium between the Earth's lithosphere and asthenosphere such that the tectonic plates (continental and ocean crusts) "float" at an elevation which depends on their thickness and density. (similar to ice floating in water).

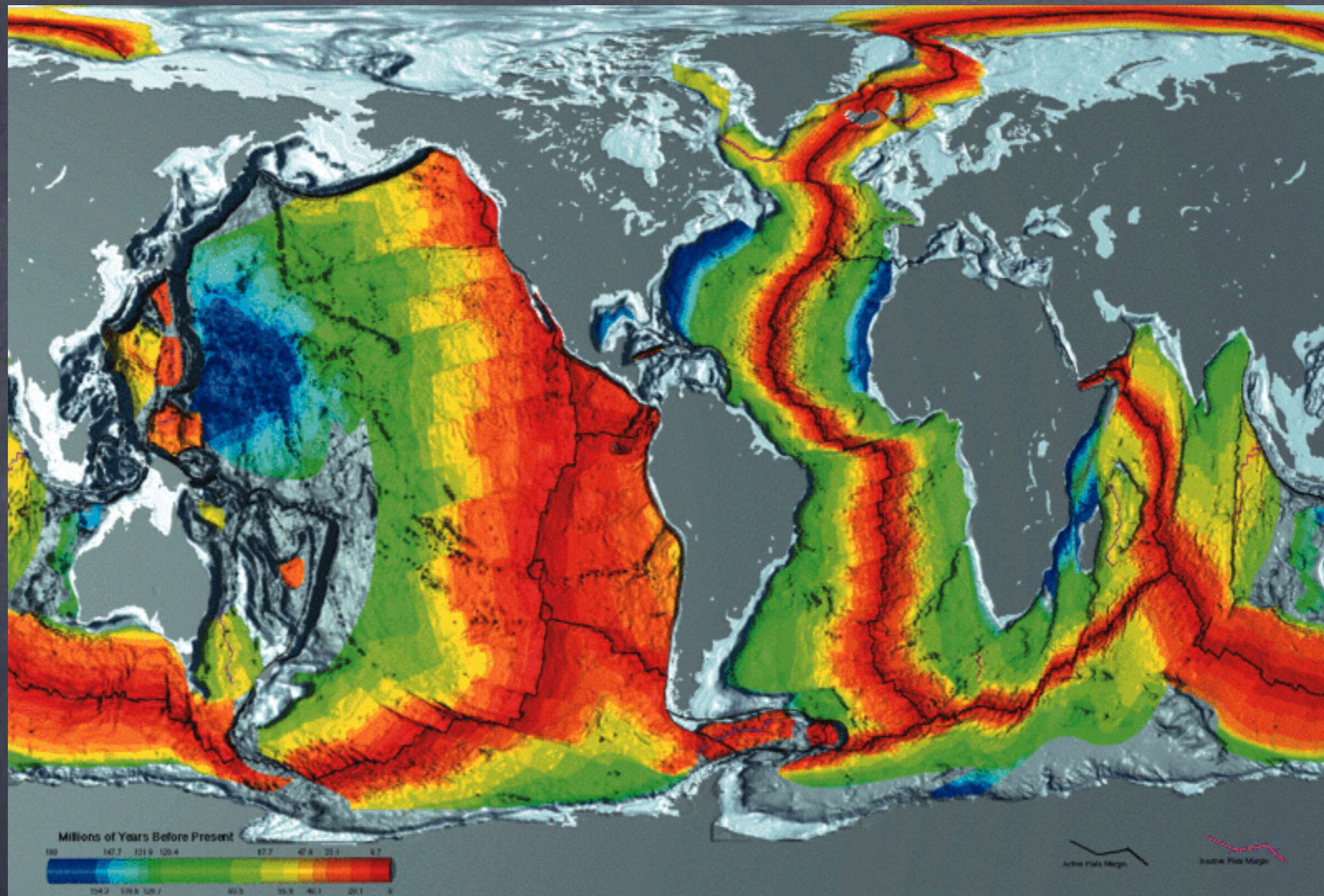


In the 1960's there was a geological revolution: the realization that the surface of the earth is in motion, slowly recycling the material that makes up our environment and shapes the ocean basins and seafloors.



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*The major lithospheric plates*

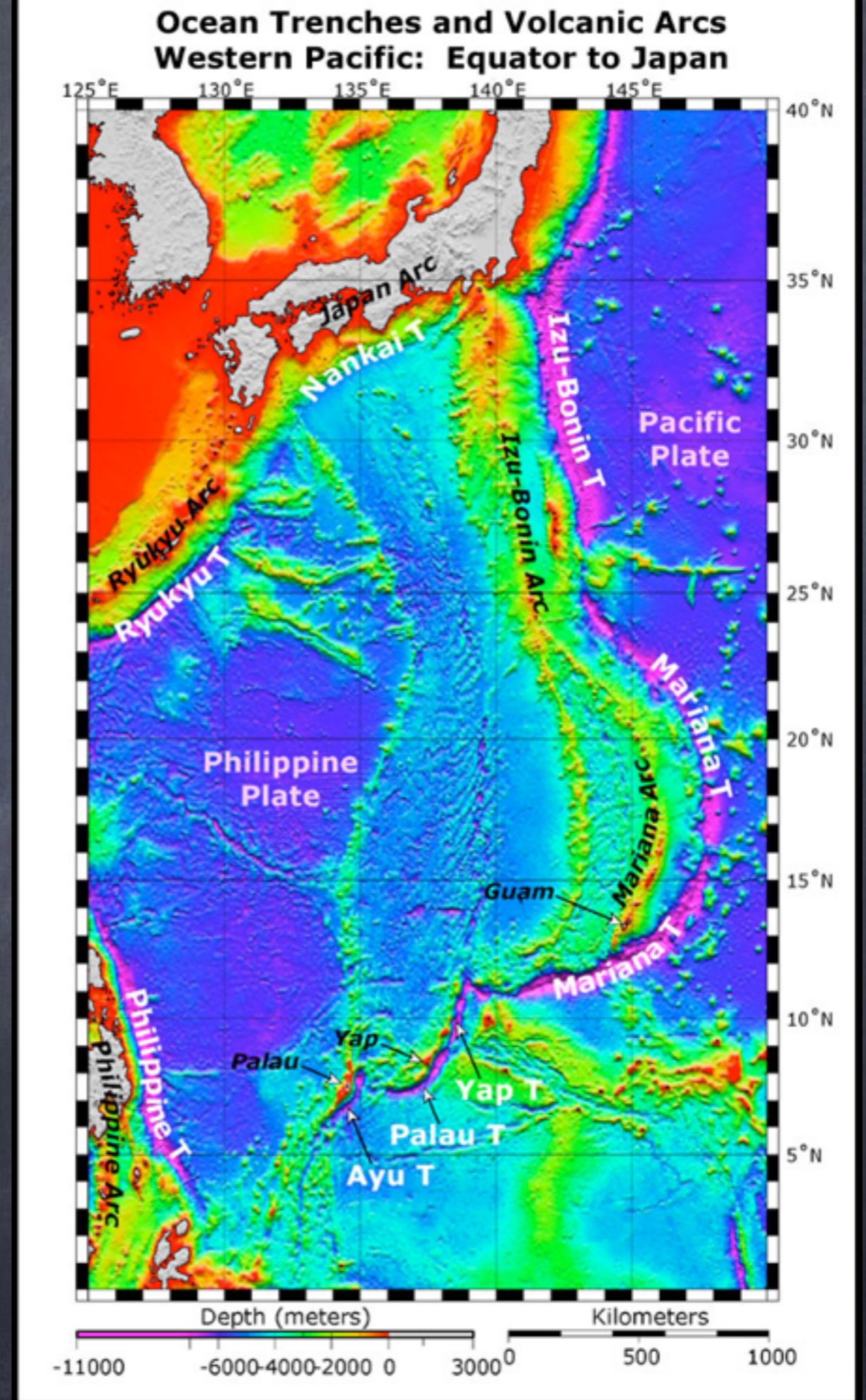


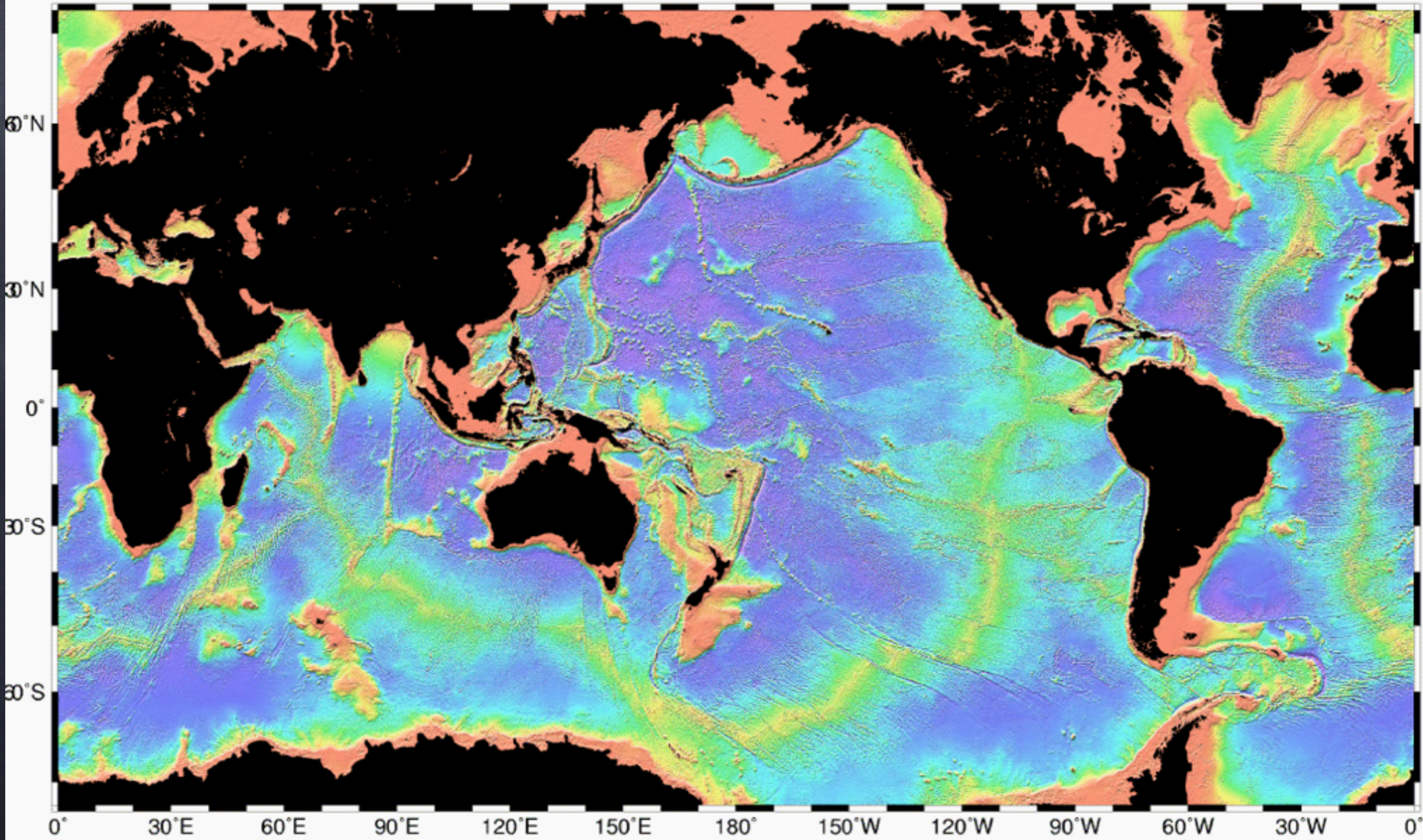
Age of oceanic crust.  
Youngest crust is along spreading centers (in red) – these are the mid-ocean ridges

Satellite altimetry data of the western Pacific from the equator to Japan. Submarine trenches are magenta. Volcanic Arcs are indicated by italicized text.

In the western Pacific the volcanic arcs (island and submarine) are west of the **deep ocean trenches**, due to the western migration and subduction of the Pacific plate under the Philippine plate. Volcanic arcs form when the subducting plate melts at depth and magma is produced, which rises buoyantly and forms volcanic arc lavas.

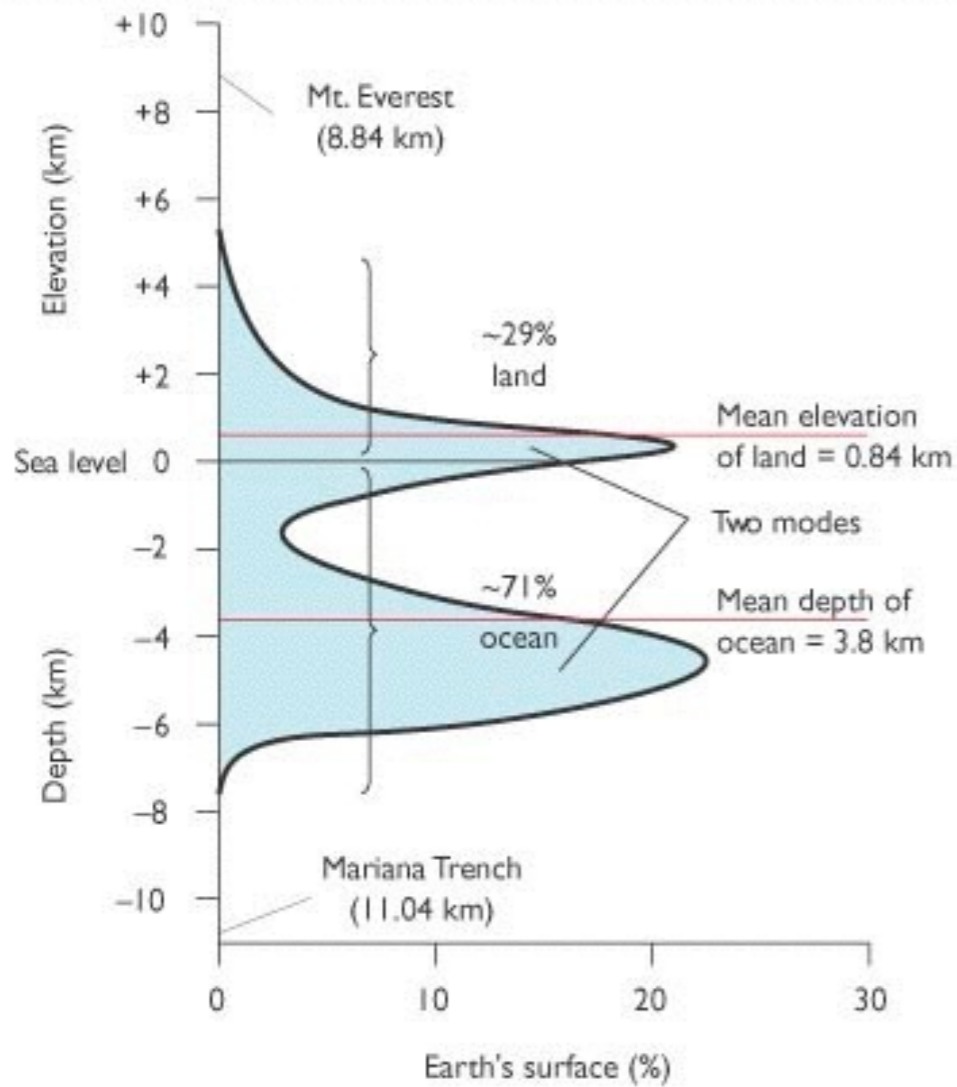
Satellite data courtesy of Smith and Sandwell, 1997. Image courtesy of the NOAA Vents Program.



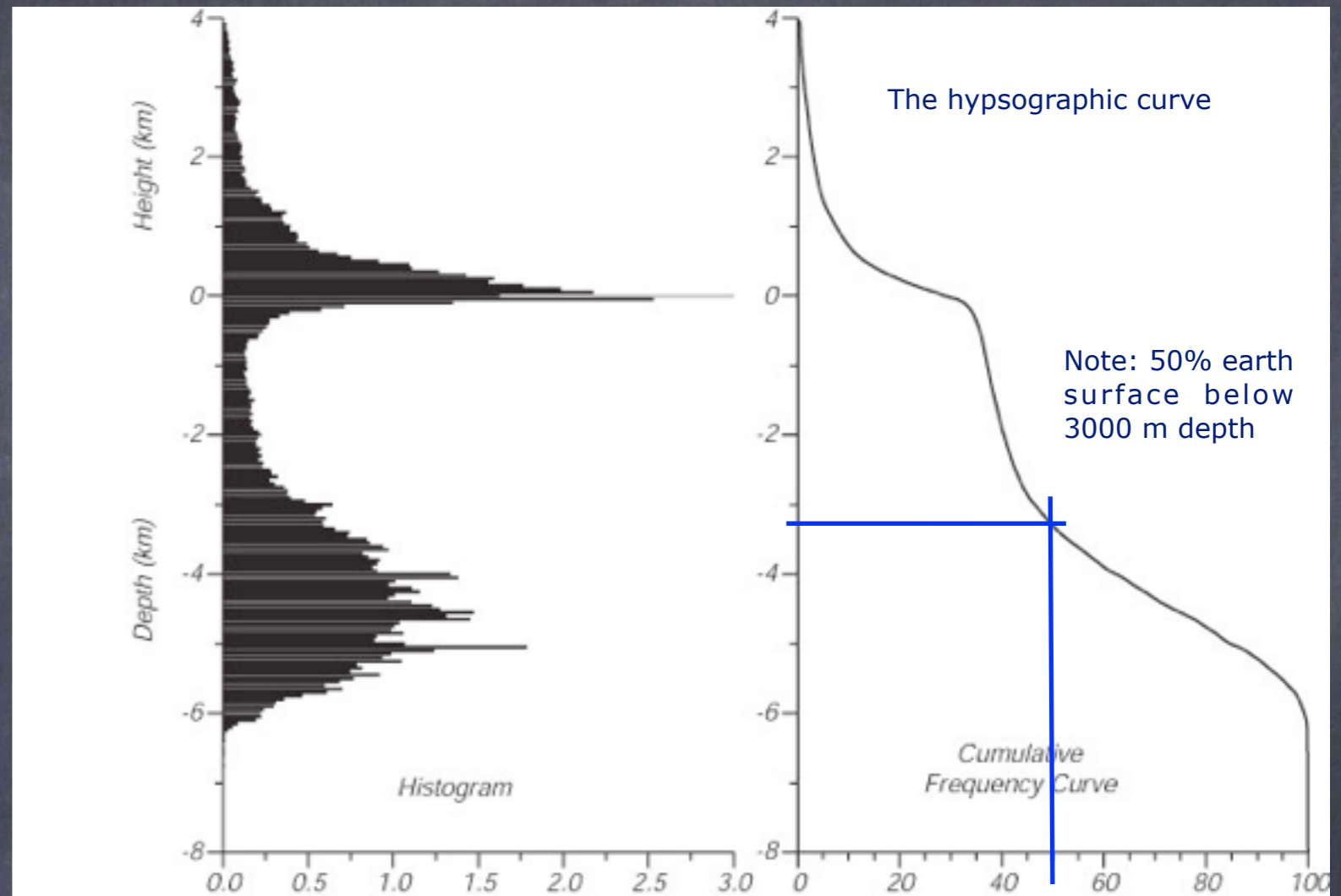


Walter H. F. Smith and David T. Sandwell, Seafloor Topography Version 4.0, SIO, September 26, 1996

Copyright 1996, Walter H. F. Smith and David T. Sandwell



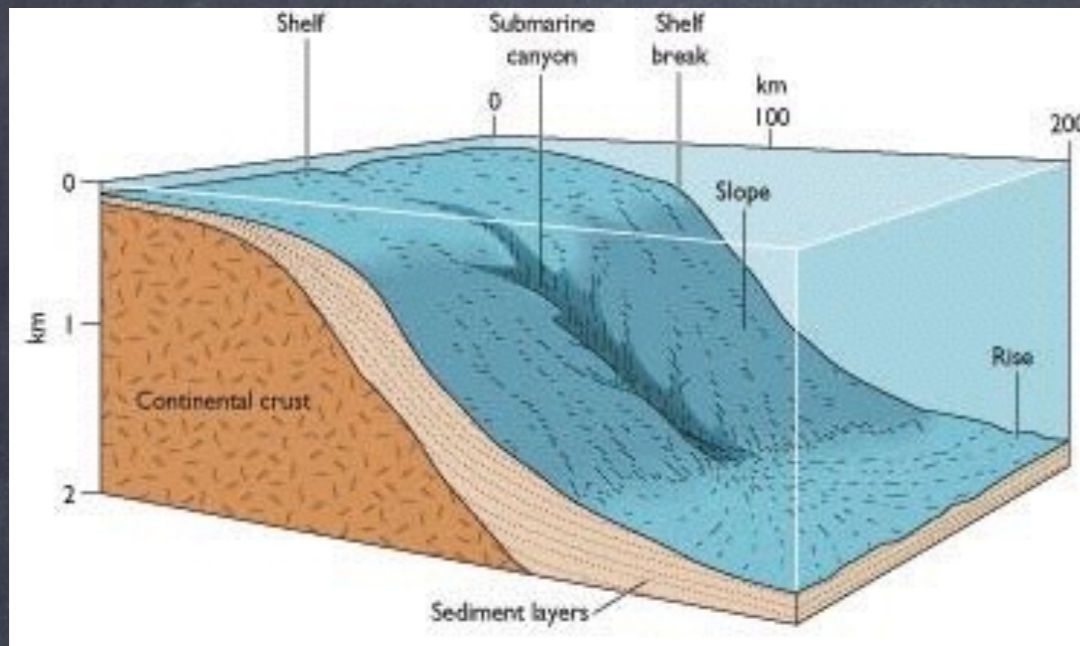
(b) FREQUENCY PLOT OF TOPOGRAPHY AND BATHYMETRY



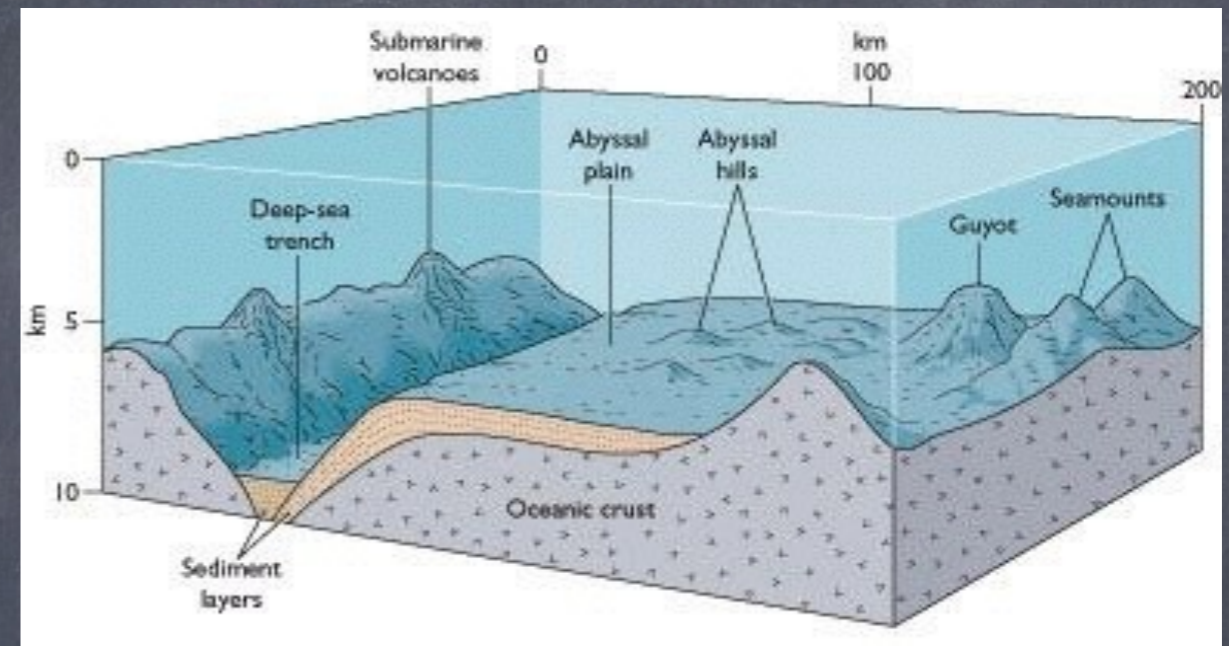
%

Elevation of Earth's surface displays a bimodal distribution with about **29% above sea level** and much of the remainder at a depth of 4 to 5 kilometers below sea level.

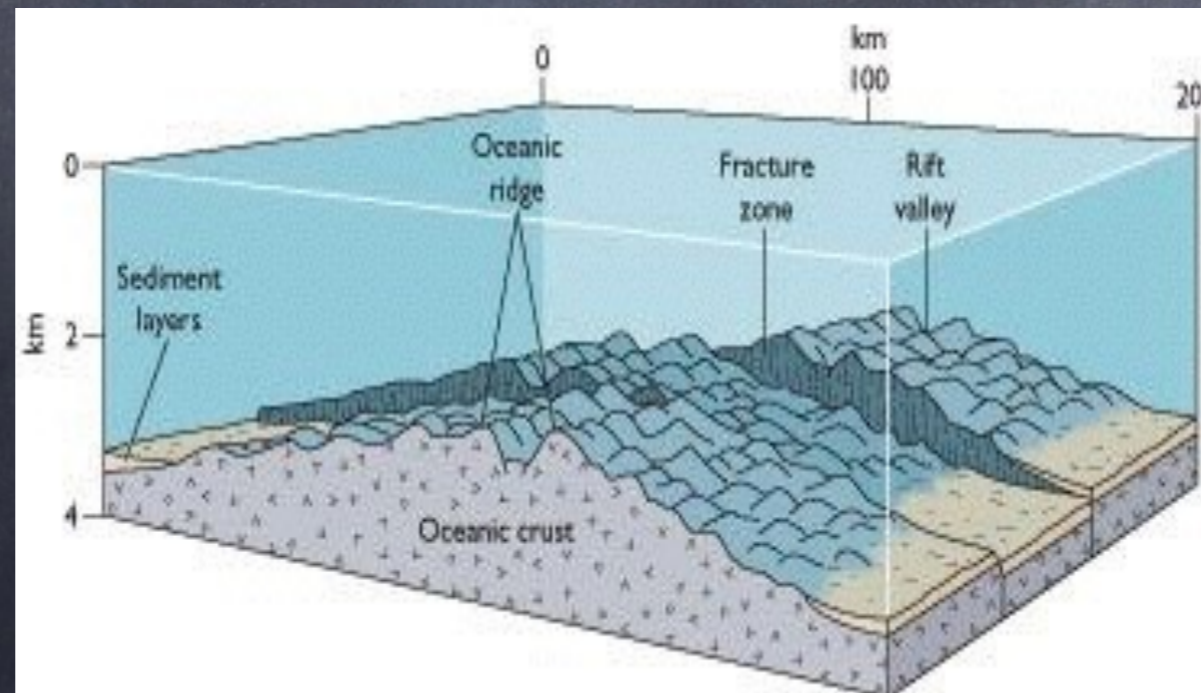
# Continental margin

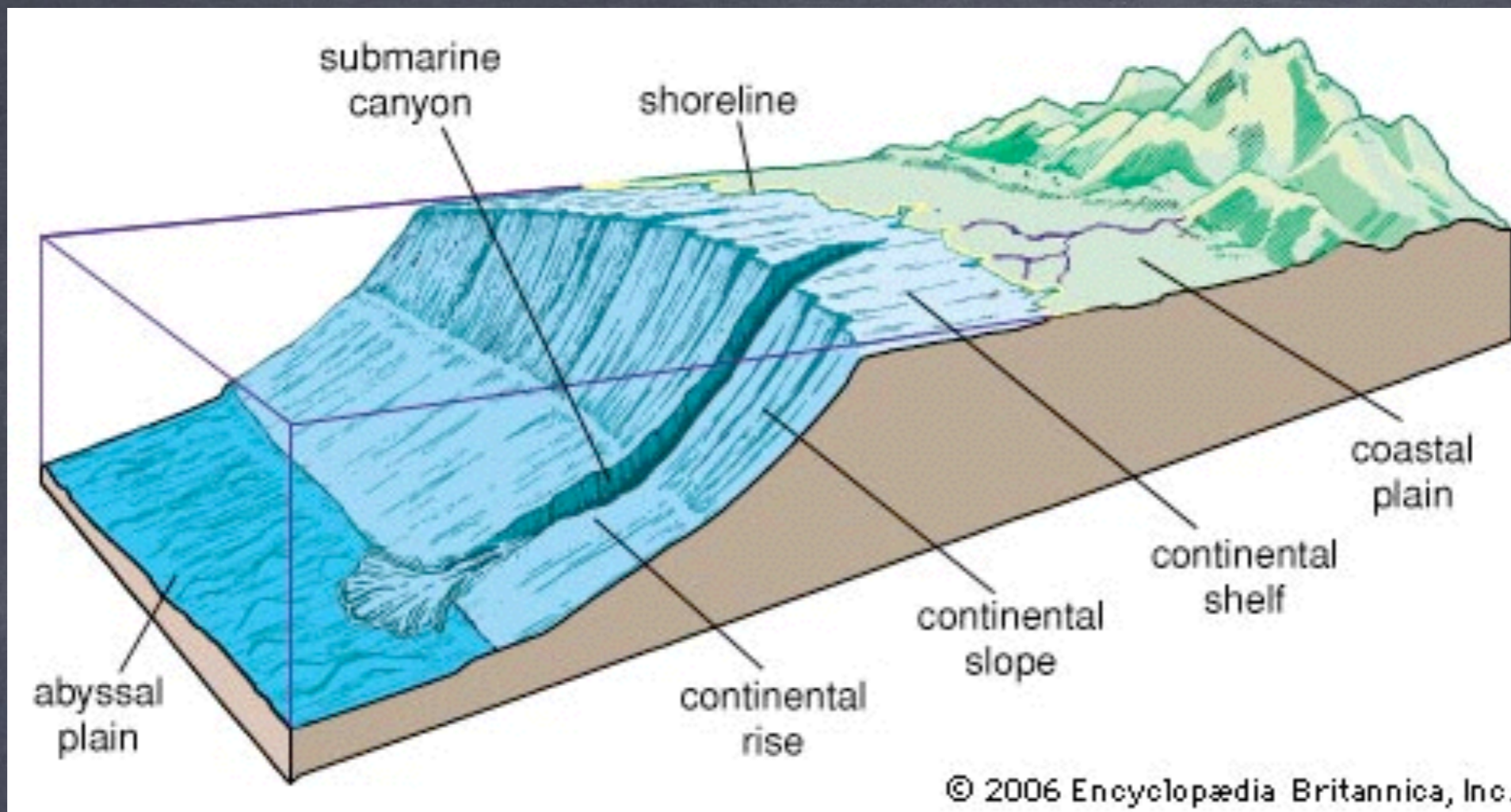


# Ocean basin



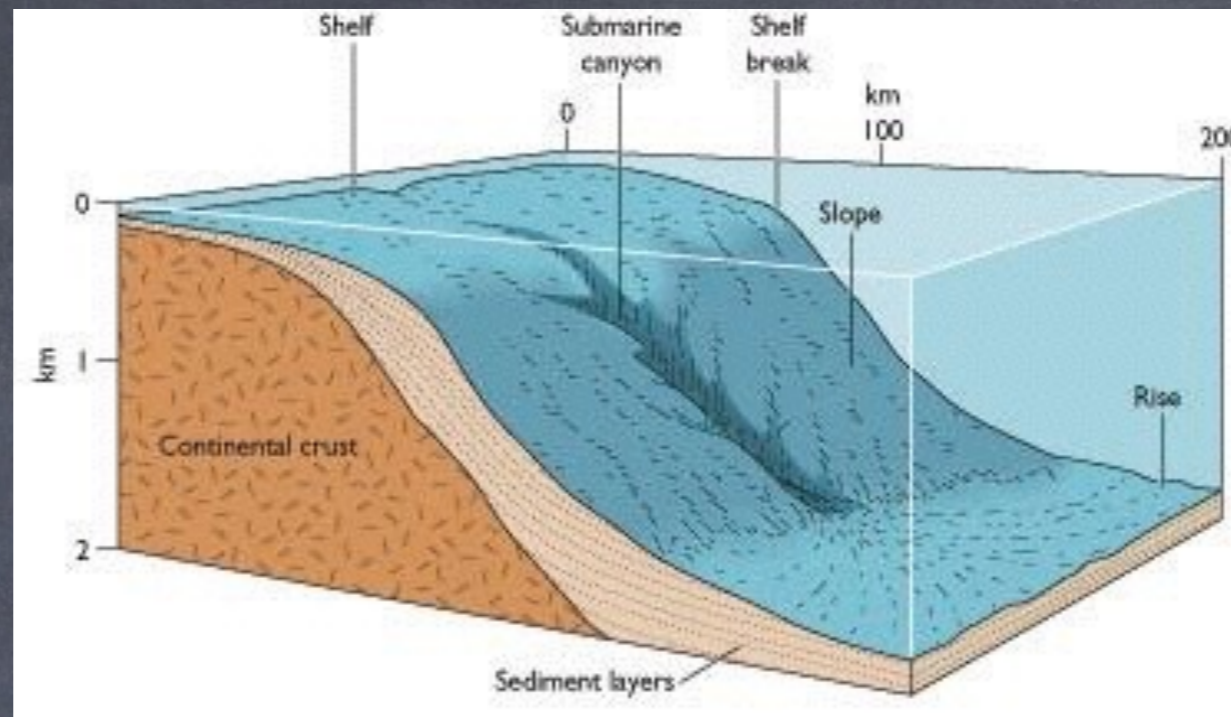
# Mid-ocean ridge





Continental margins are the submerged edges of the continents and consist of massive wedges of sediment eroded from the land and deposited along the continental edge. The continental margin can be divided into three parts:

Continental shelf,  
Continental slope,  
Continental rise.



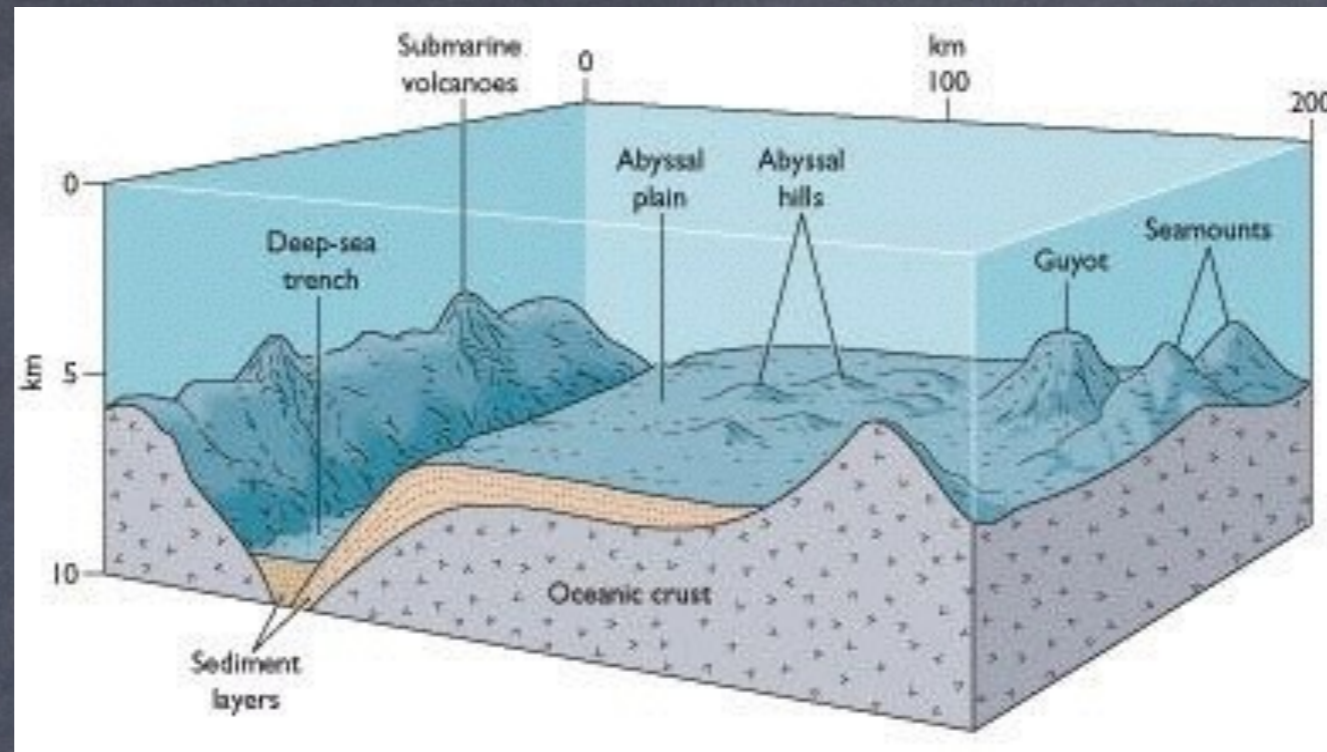
Feature	Width	Water Depth	Bottom Gradient
Shelf	1-1000km	<150m	<1:1000 ( $\sim 0.5^\circ$ )
Slope	10-200km	100m to several km	$\sim 1:40$ ( $4^\circ$ - $6^\circ$ )
Rise	100-500km	3000-4000km	$\sim 1:700$ ( $0.5^\circ$ - $1^\circ$ )

20,000 years ago the continental shelves were above sea level.

Submarine canyons may have been carved by rivers or by gravity/turbidity currents.



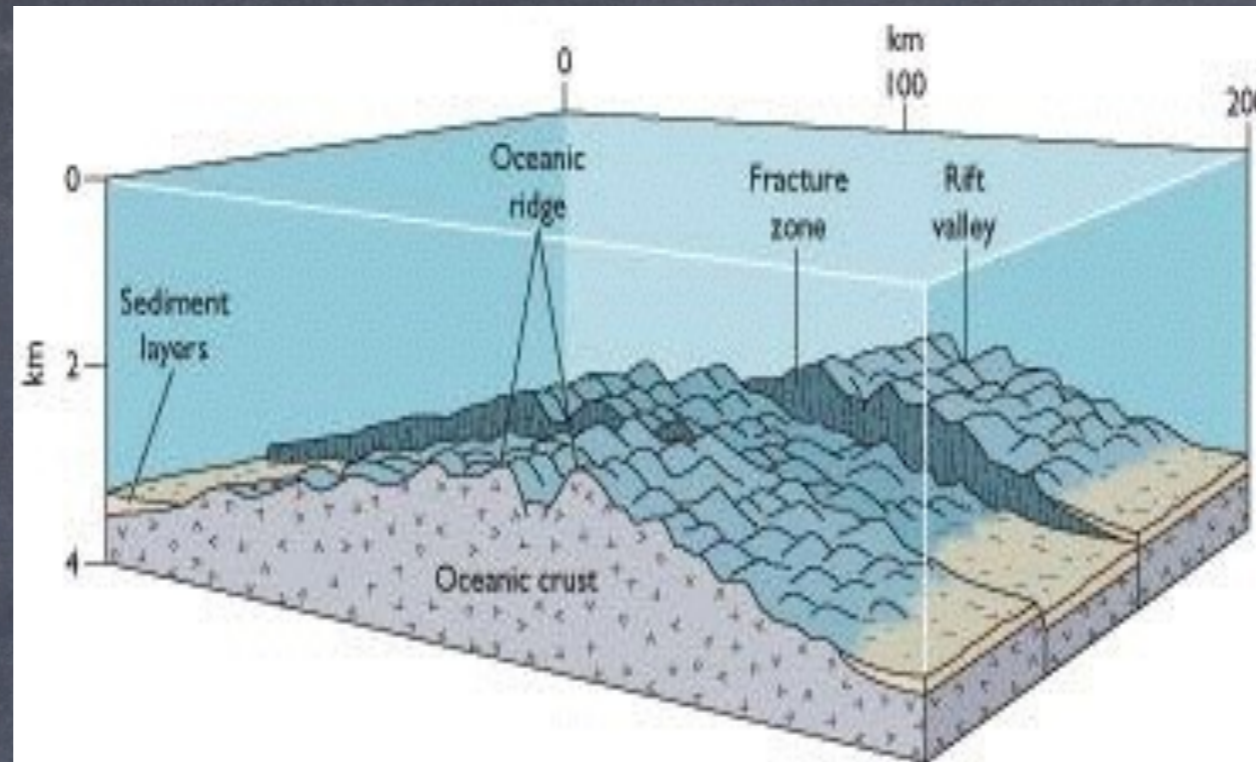
# Ocean basin



Feature	Width	Water Depth	Bottom Gradient
Abyssal plains	1-1000km	> 3km	<1:1000 (<0.5°)
Abyssal hills	0.1-100km	variable	
Seamounts	2-100km	variable	
Deep-sea trench	30-100km	5-12km	

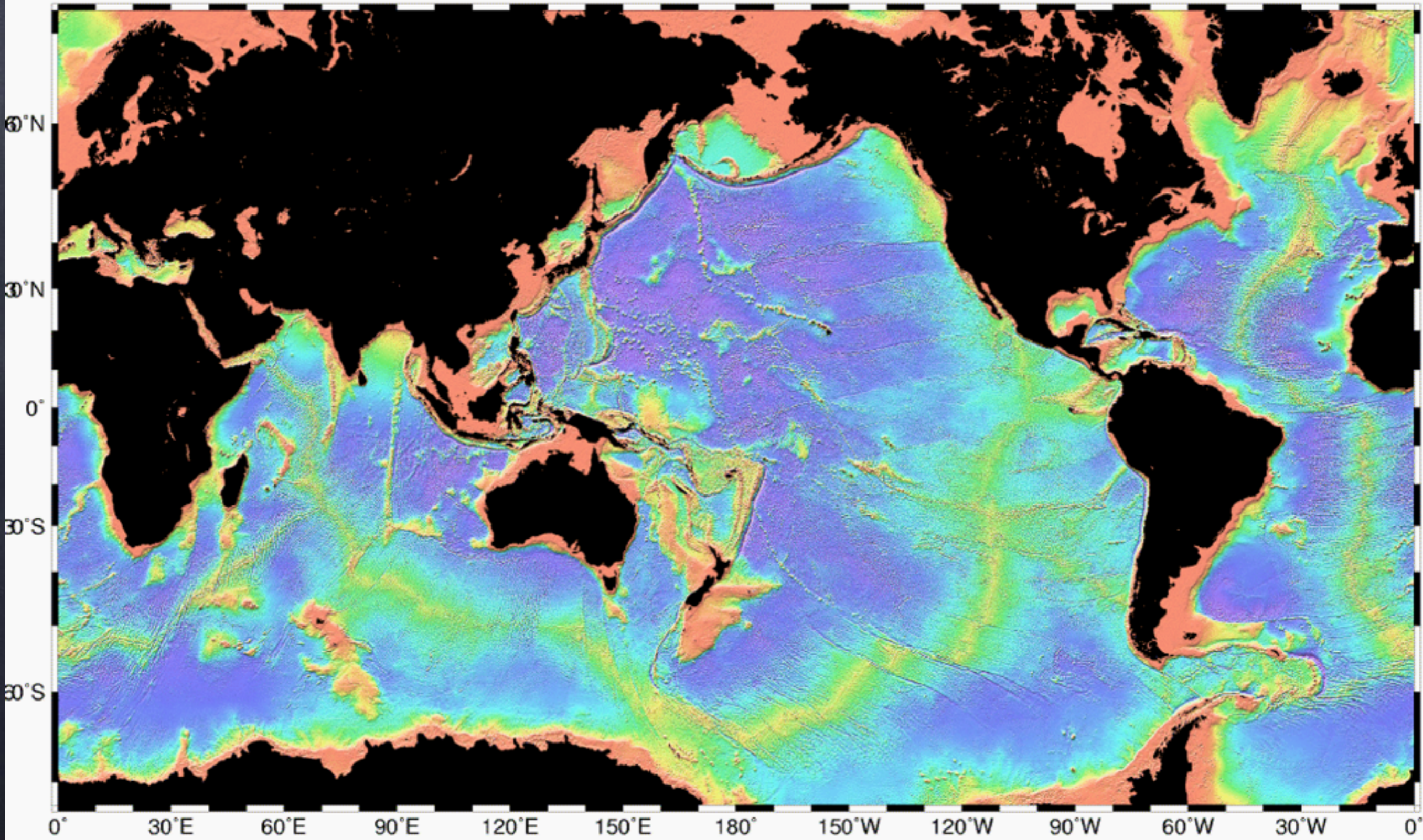
Seamounts can reach to within a few hundred meters of the surface and support vibrant ecosystems, fed by upwelling waters

# Mid-ocean ridge



Feature	Width	Water Depth
Midocean ridge flanks	500-1500km	> 3km
Midocean ridge crest	500-1000km	2-4km

The mid-ocean ridge system is a continuous submarine mountain range covering about one third of the deep ocean and extending for 60,000 km around the Earth (almost 10X radius of earth).



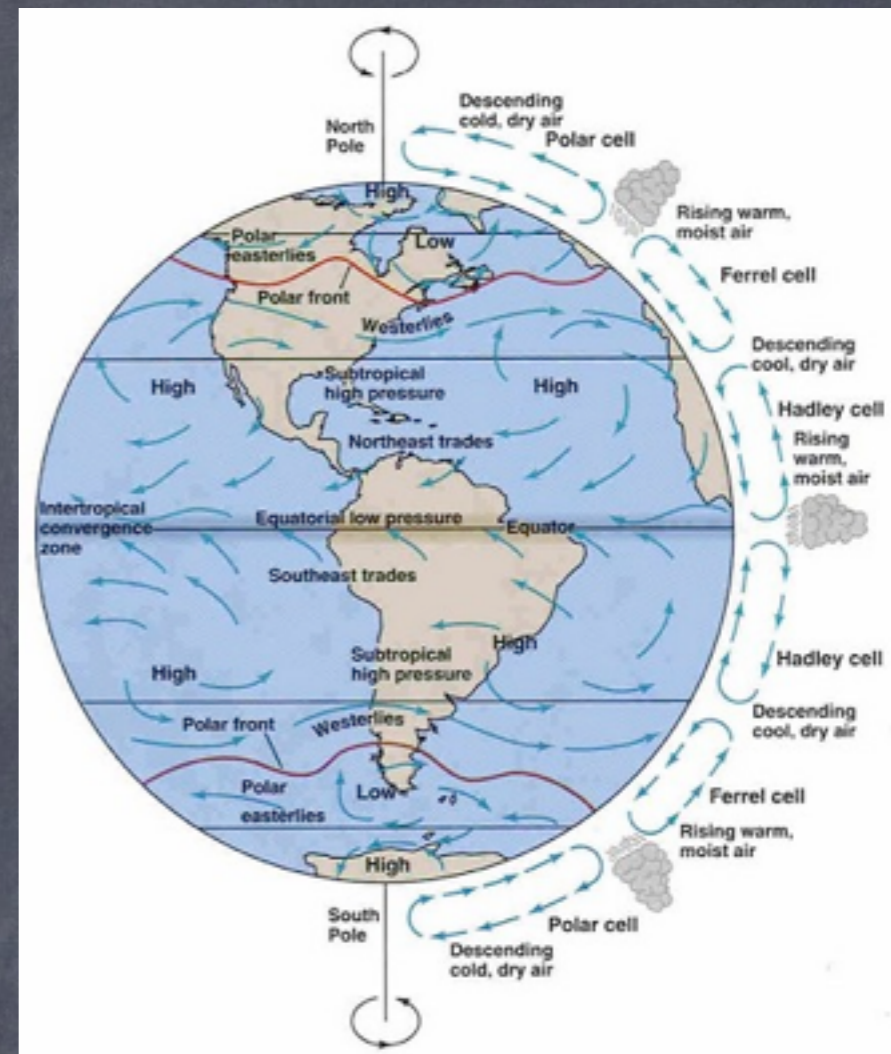
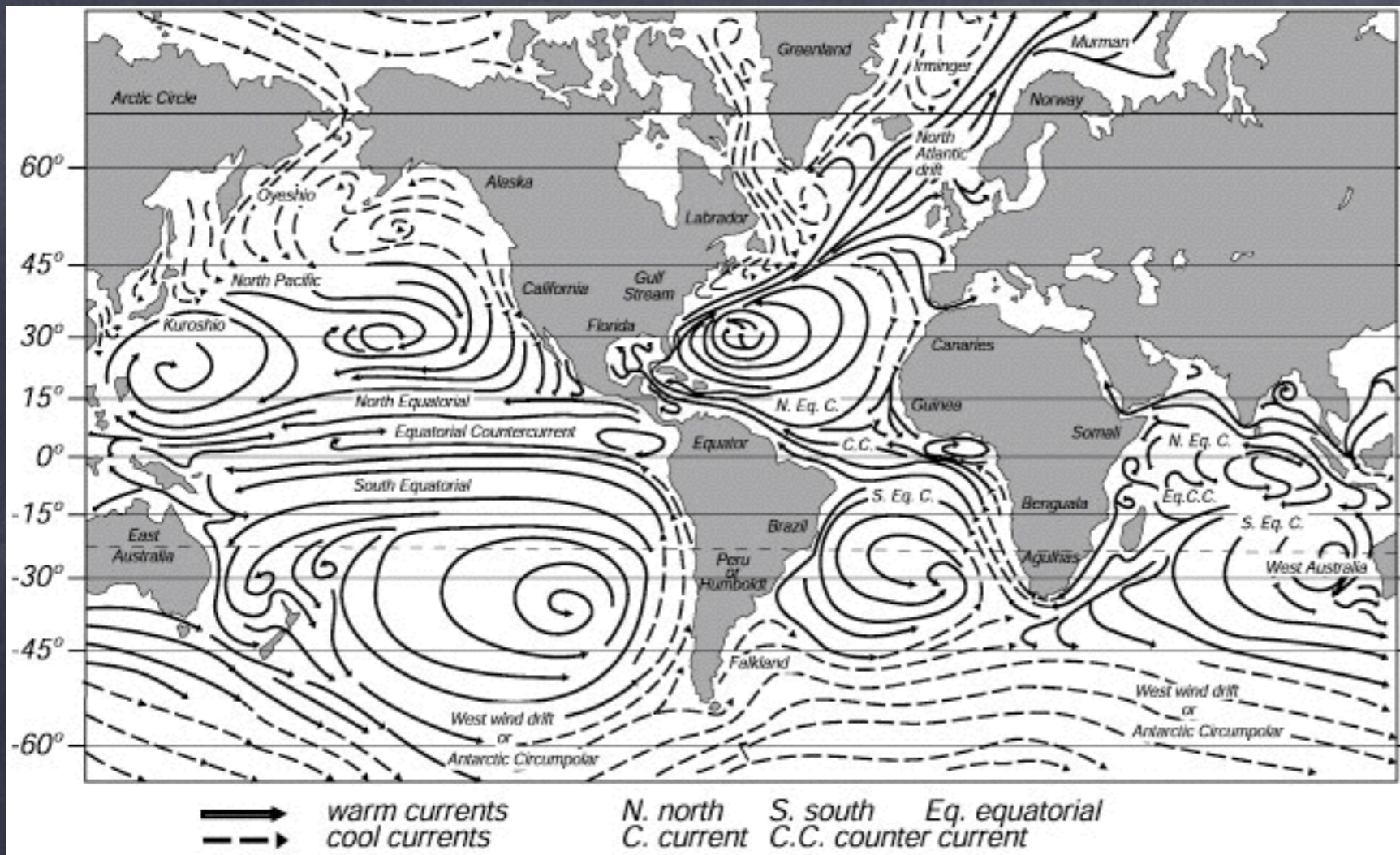
Walter H. F. Smith and David T. Sandwell, Seafloor Topography Version 4.0, SIO, September 26, 1996

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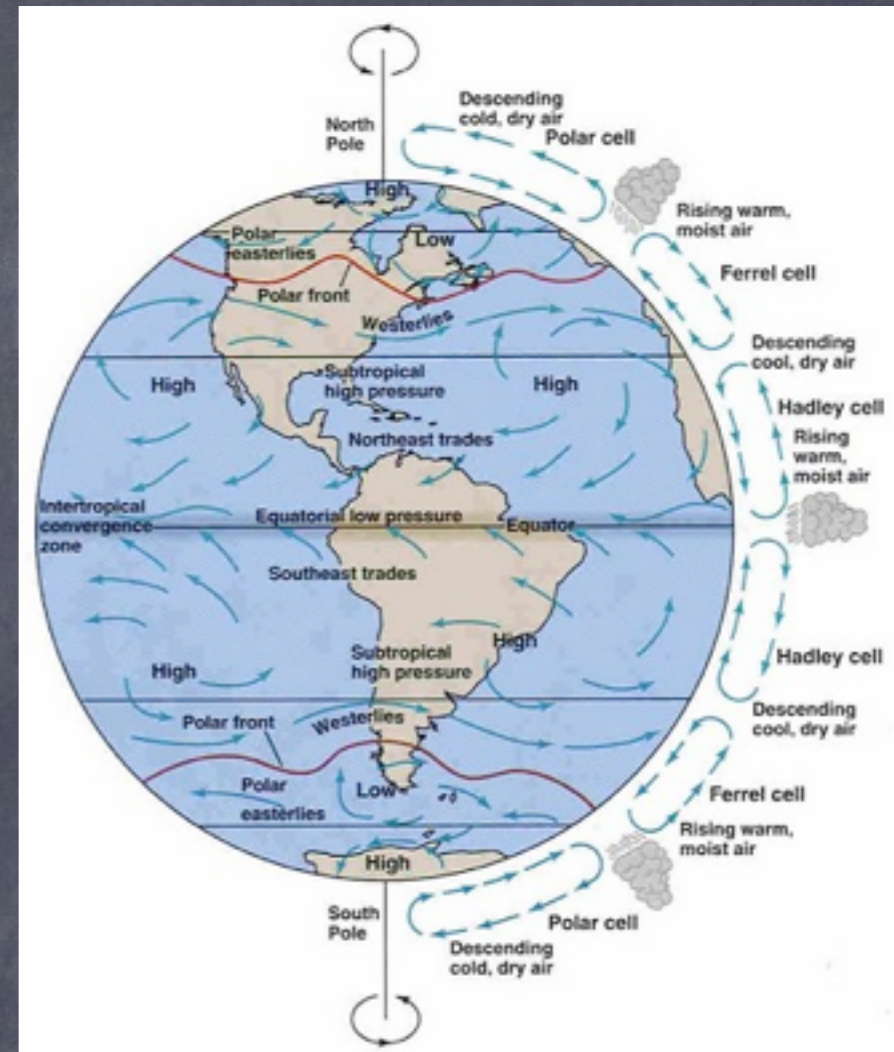
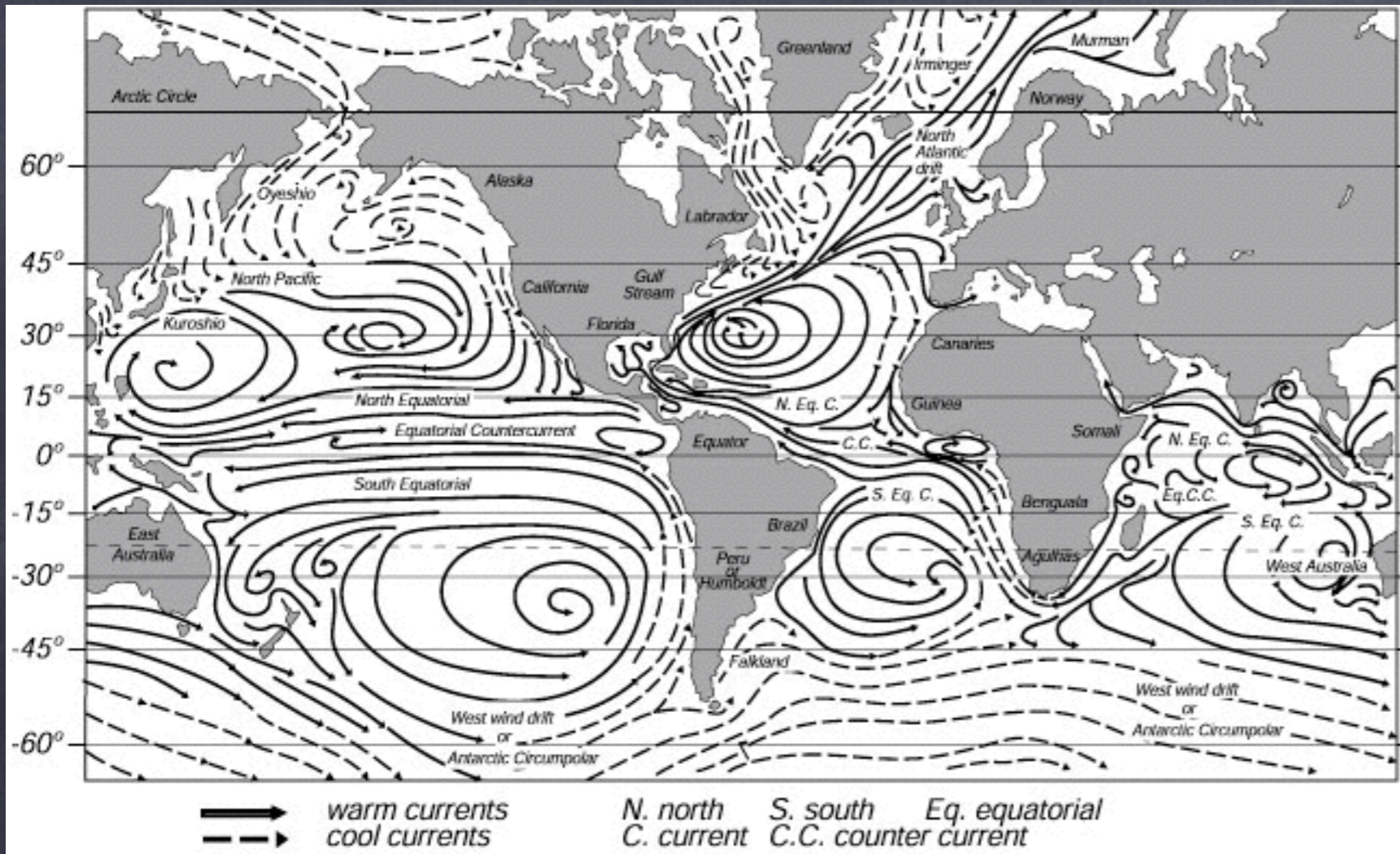
# Major current systems



You see **gyres** separated by zonal flow at equator

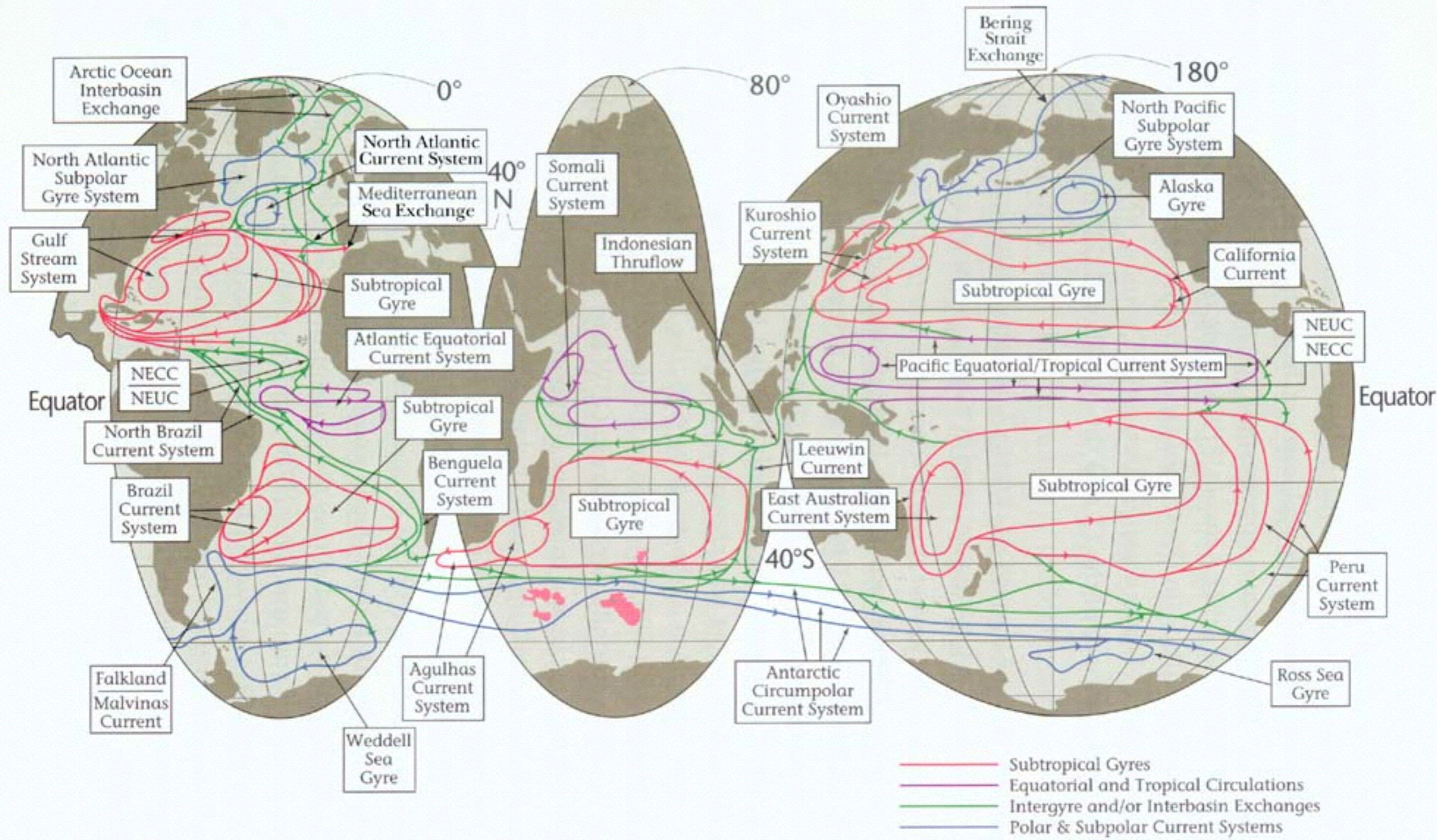


Now see gyres offset to west and zonal currents and countercurrents on equator, plus subpolar gyres in northern hemisphere, and north-south excursion of ACC in Southern Ocean



Wind systems look somewhat similar to ocean gyres, but not asymmetric... and ECC in opposite direction!

Notice: subtropical gyre centers at latitude 30 - where air descends (high P). Gyre boundaries at 45 - rising air (low P)



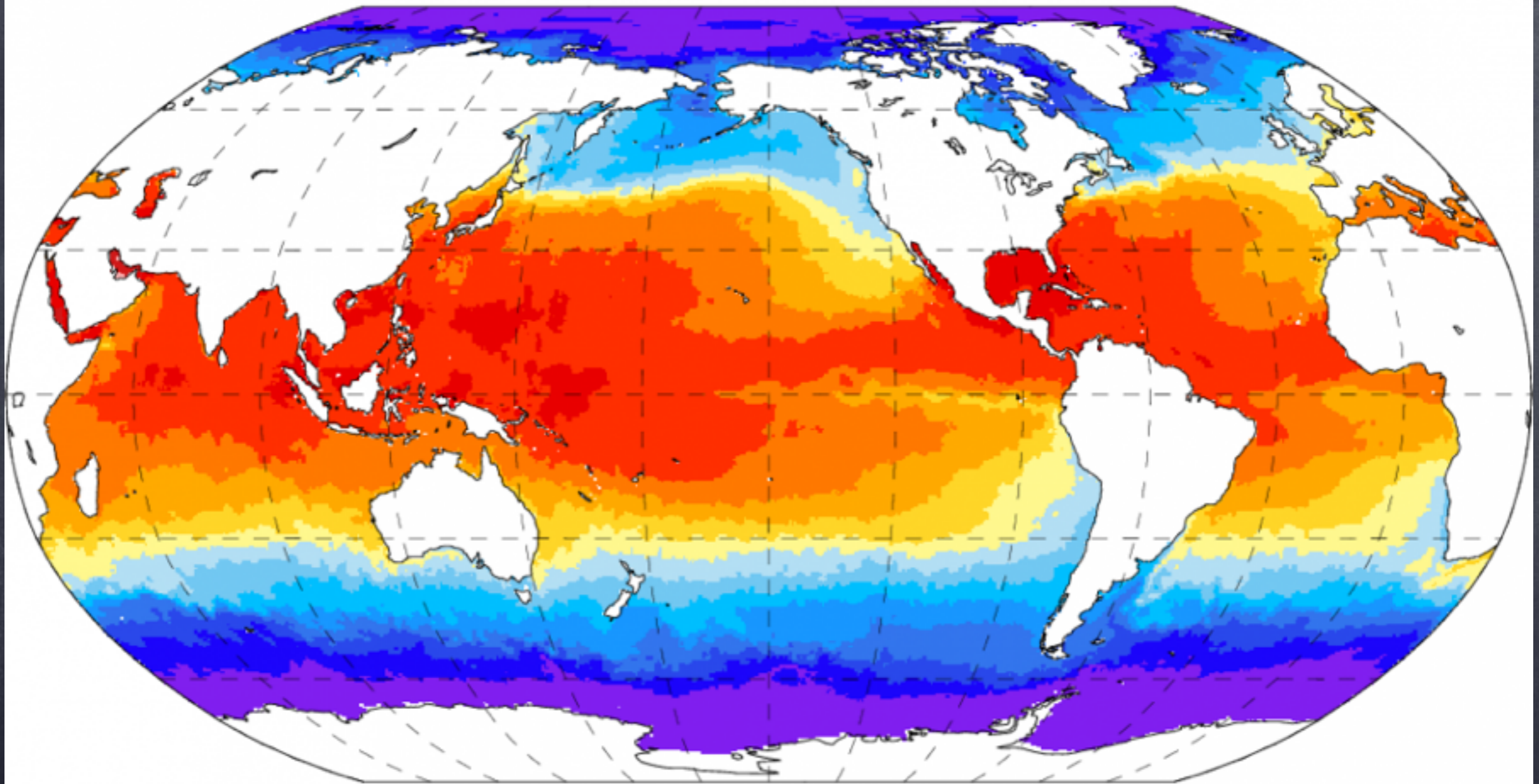


# Sea Surface Temperature

**GHRSSST: 20110806-UKMO-L4HRfnd-GLOB-v01-fv02-OSTIA**

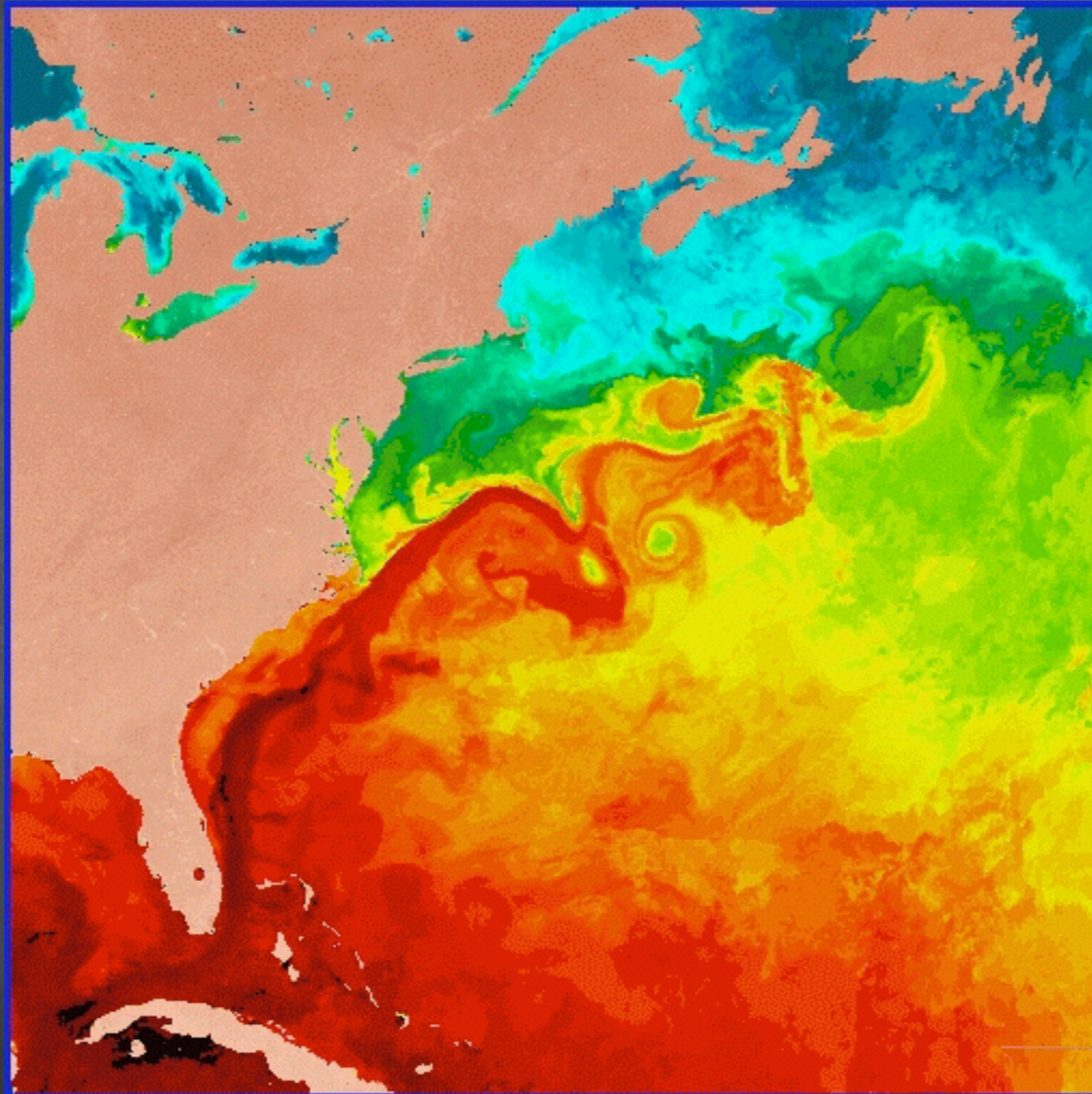
analysed sea surface temperature

kelvin



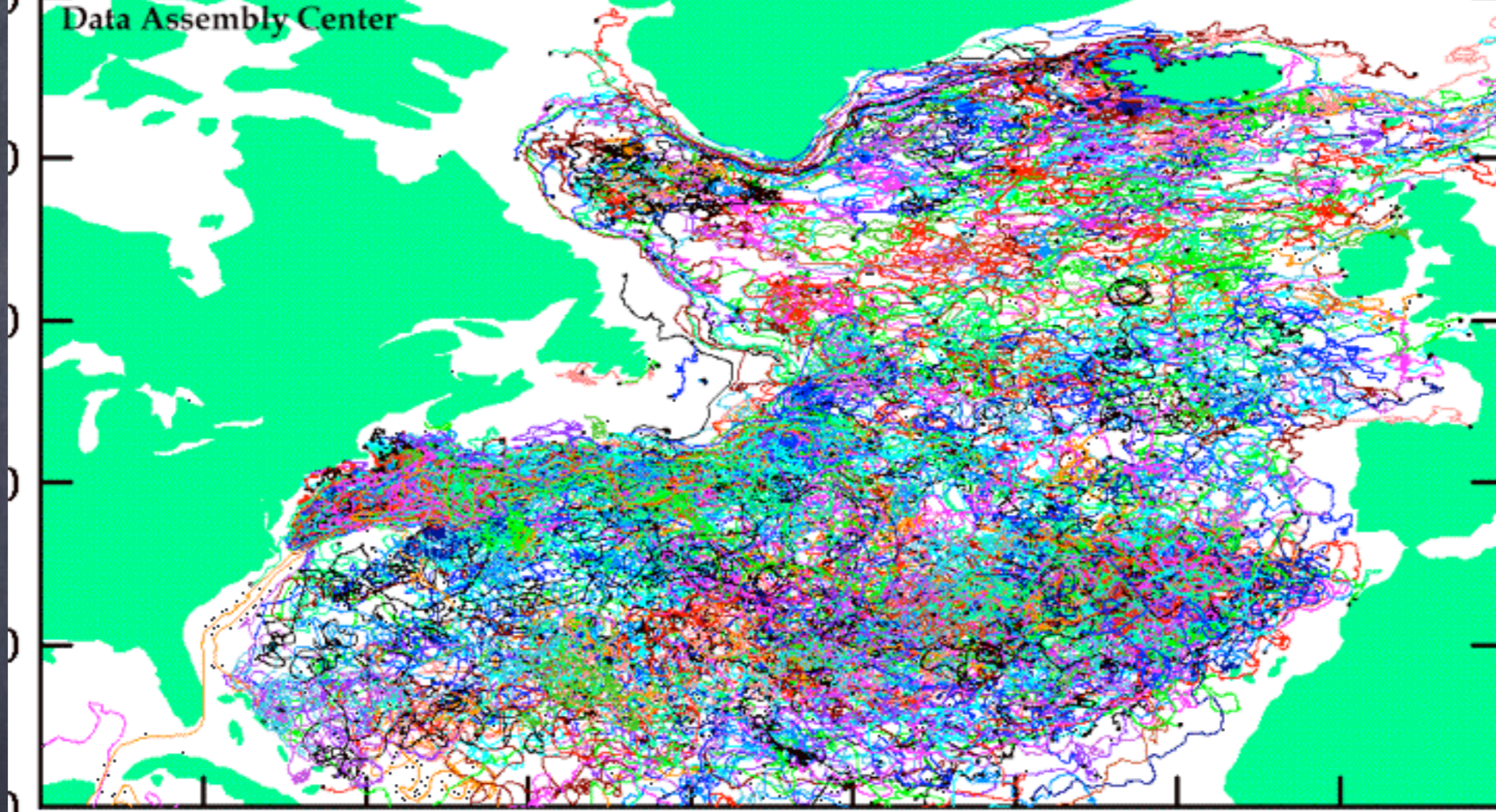
272.5 275 277.5 280 282.5 285 287.5 290 292.5 295 297.5 300 302.5 305 307.5

# Sea Surface Temperature

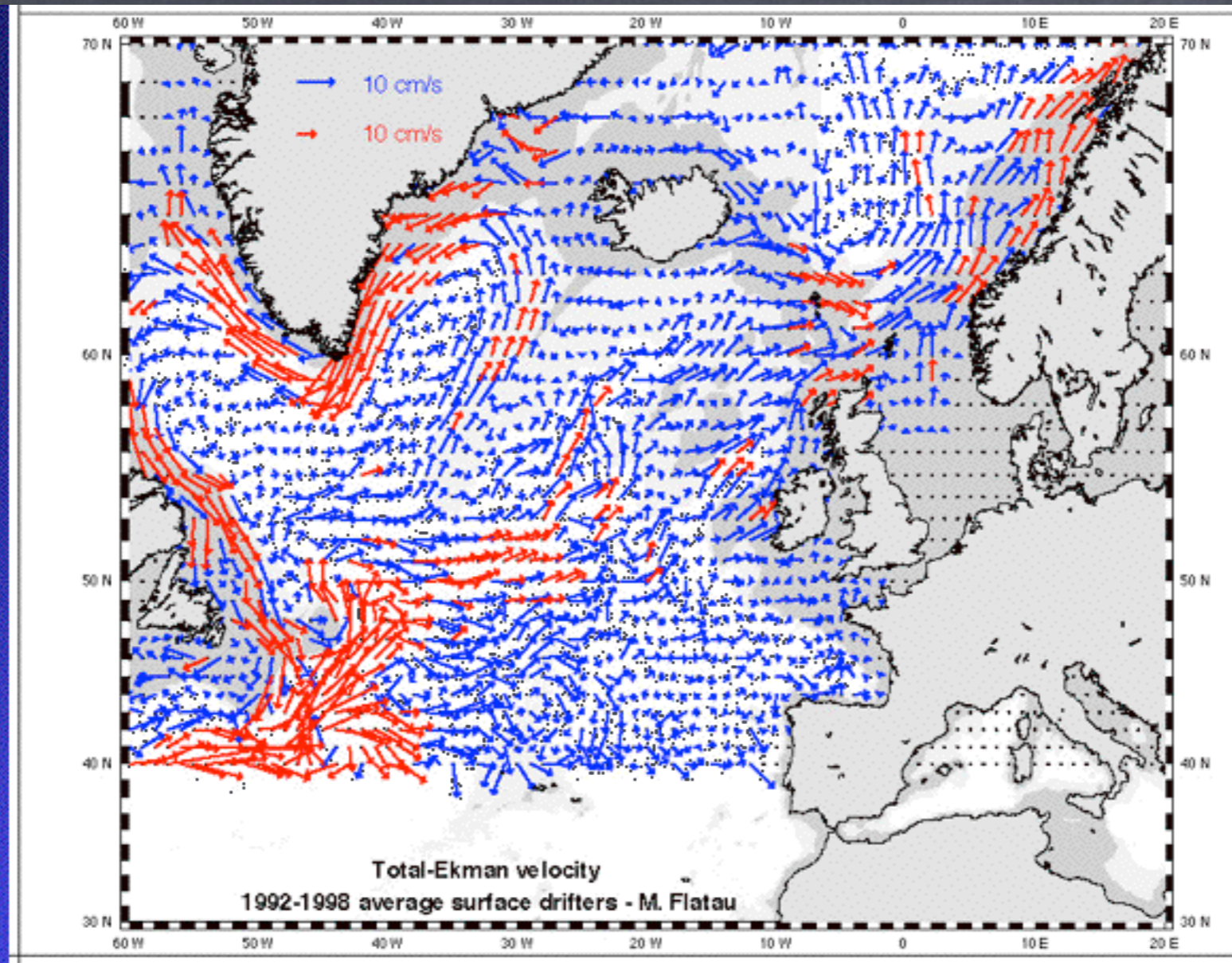


# North Atlantic Drifter Tracks Through August 1997

Data Assembly Center



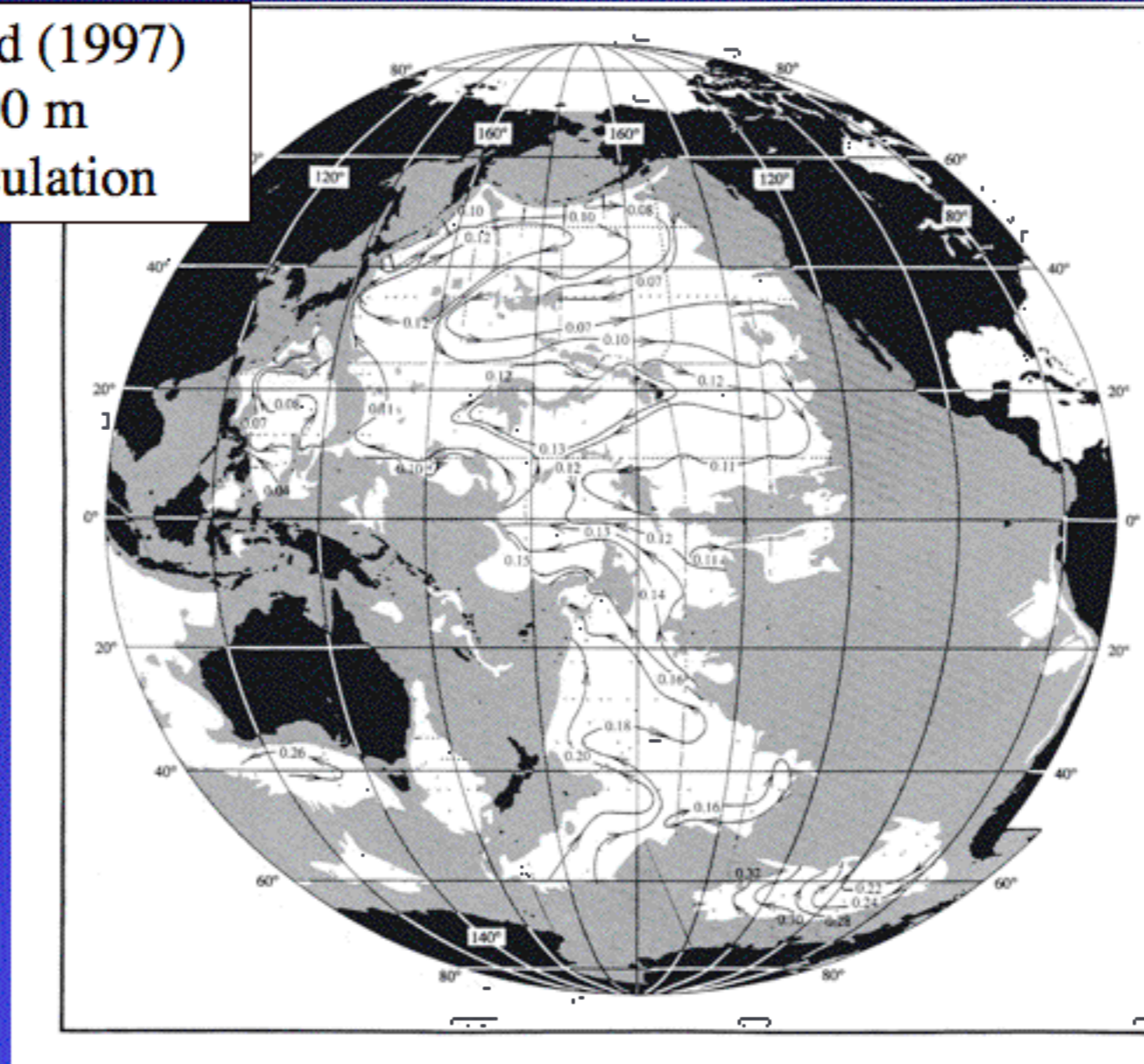
90 -80 -70 -60 -50 -40 -30 -20 -10 0



Drifter tracks averaged over six years

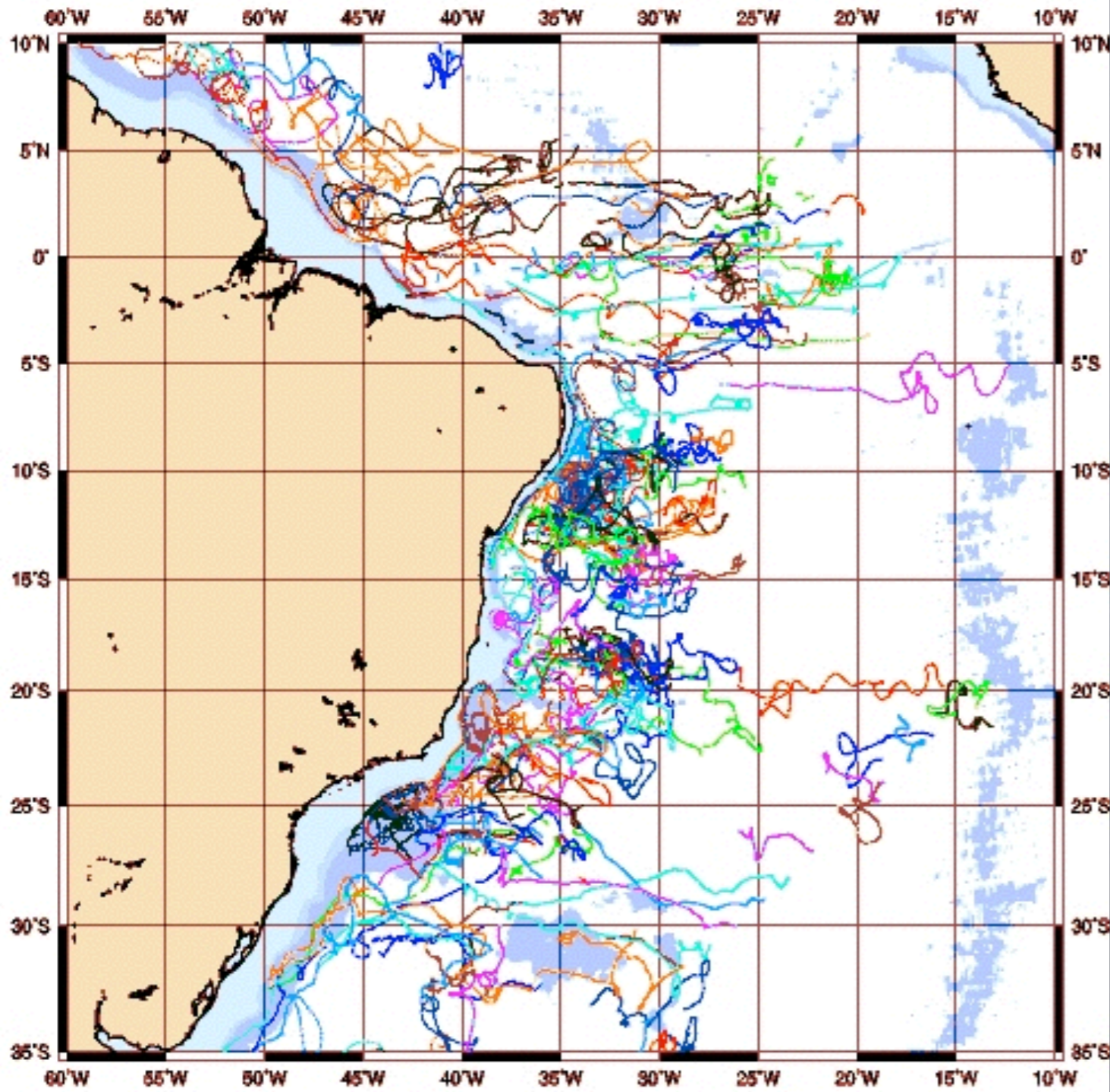
## Pacific 4500 dbar circulation

Reid (1997)  
4500 m  
circulation

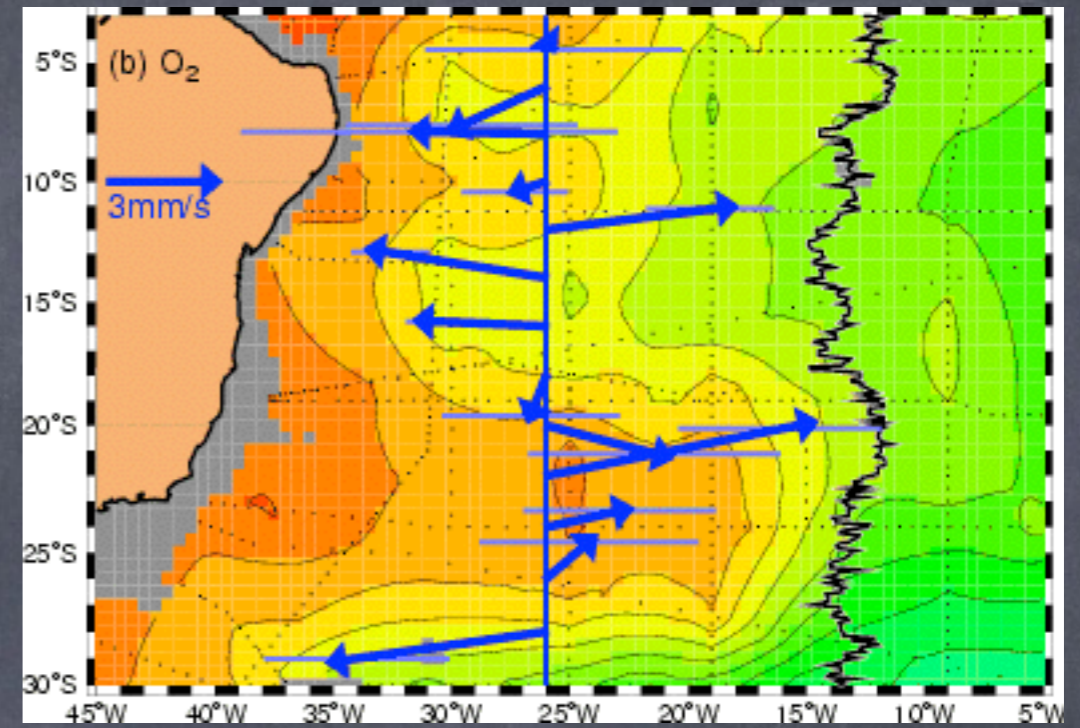


Deep circulation steered by  
bathymetry of ocean floor

750-950 dbar MARVOR, RAFOS and SOFAR float trajectories

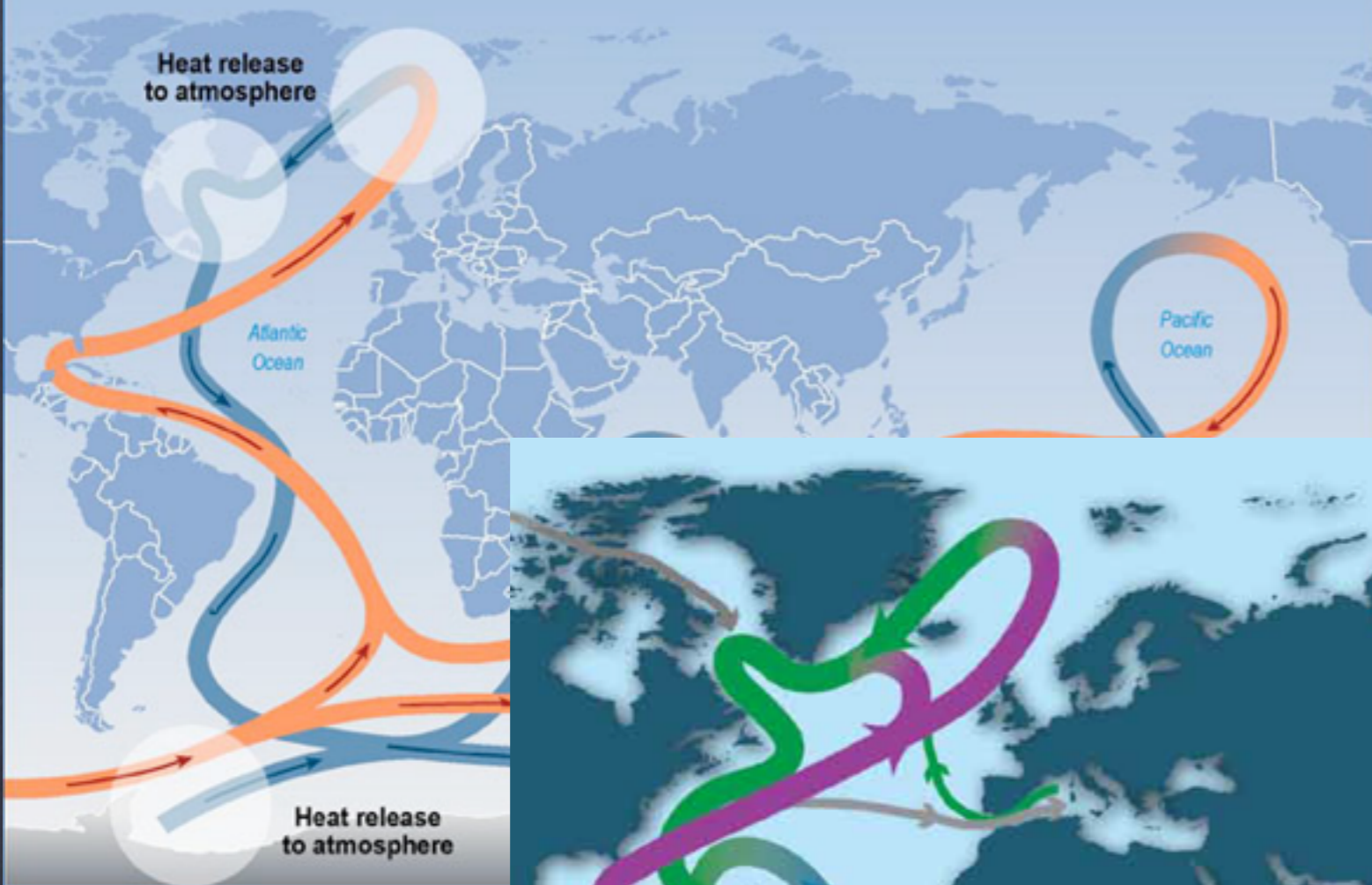


GMT 4.6.18 10:10 MercatorProjection BROWNLPO



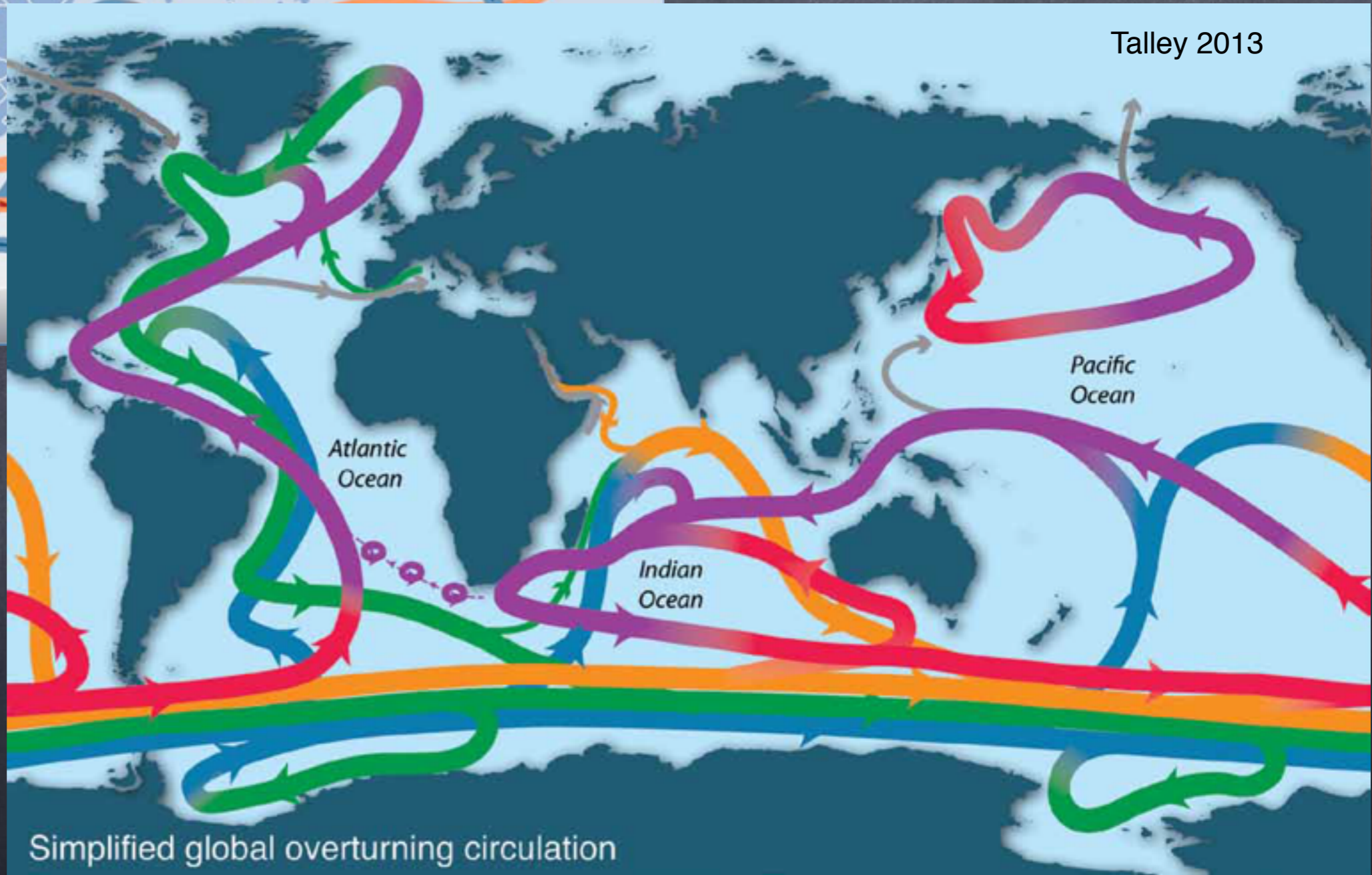
Intermediate and Deep circulation more complex and energetic than previously thought

### Great ocean conveyor belt

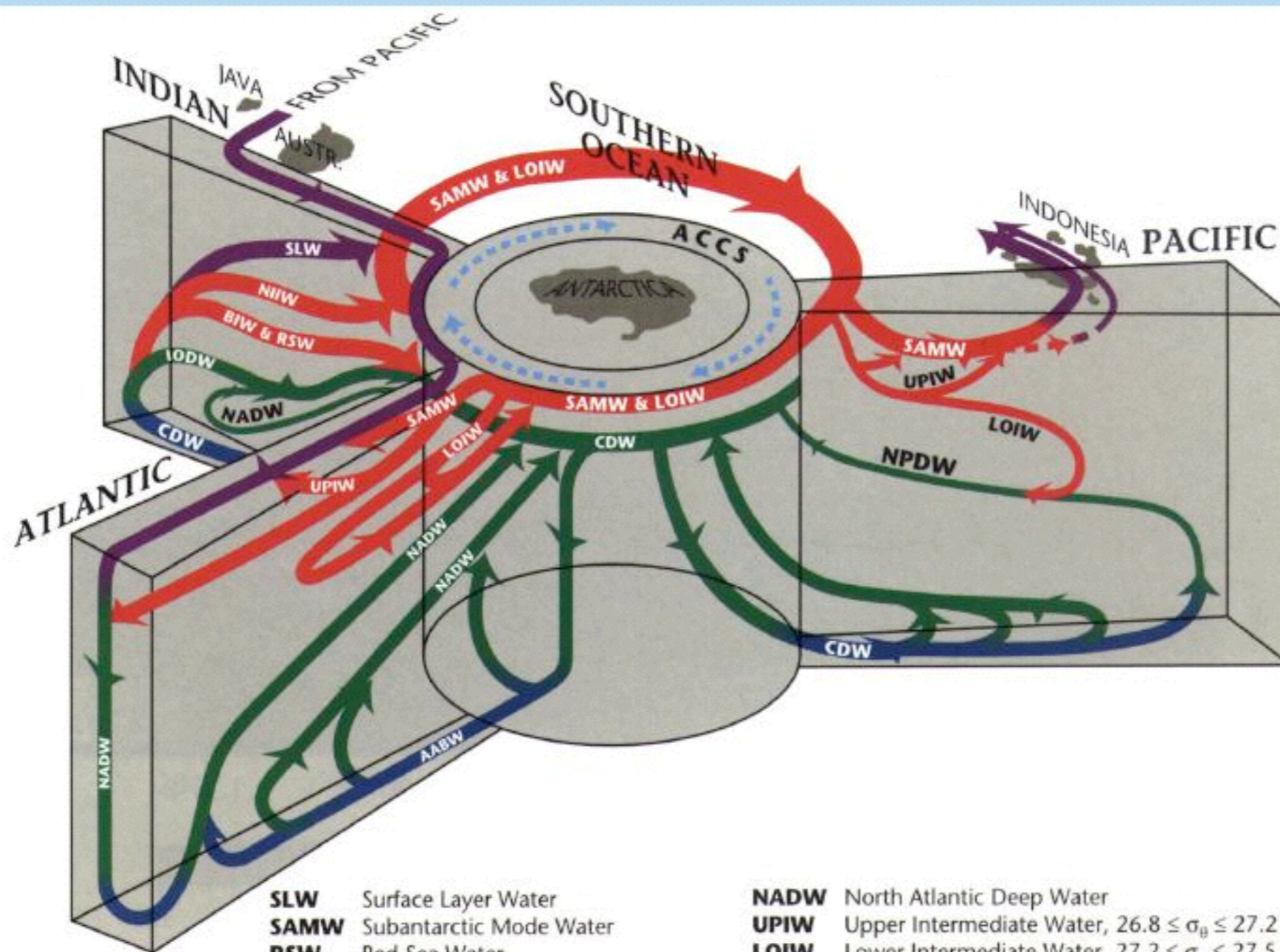


Global thermohaline circulation  
Meridional overturning circulation

Talley 2013



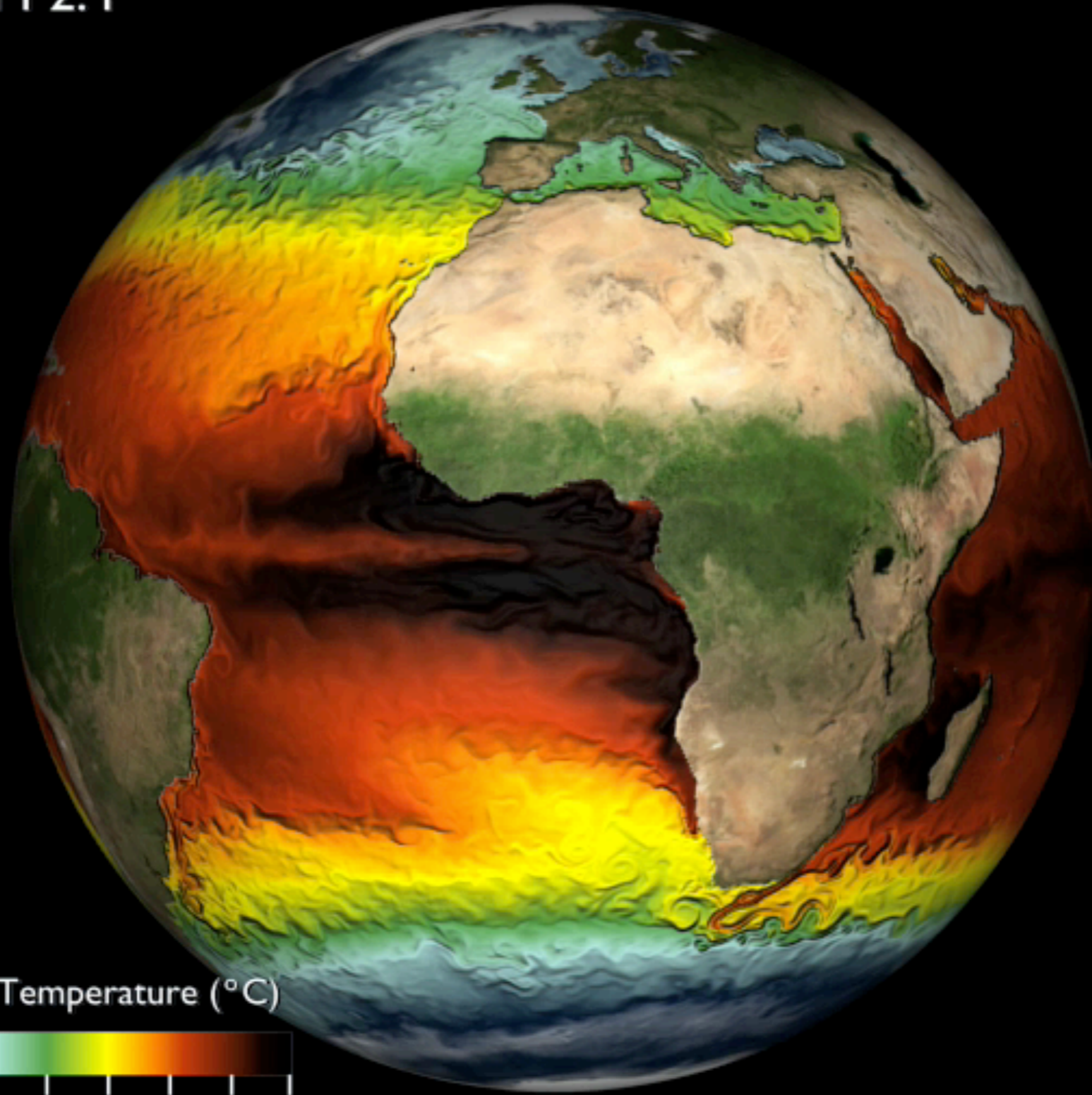
Simplified global overturning circulation



- |             |                                      |             |   |
|-------------|--------------------------------------|-------------|---|
| <b>SLW</b>  | Surface Layer Water                  | <b>NADW</b> | North Atlantic Deep Water                                       |
| <b>SAMW</b> | Subantarctic Mode Water              | <b>UPIW</b> | Upper Intermediate Water, $26.8 \leq \sigma_{\theta} \leq 27.2$ |
| <b>RSW</b>  | Red Sea Water                        | <b>LOIW</b> | Lower Intermediate Water, $27.2 \leq \sigma_{\theta} \leq 27.5$ |
| <b>AABW</b> | Antarctic Bottom Water               | <b>IODW</b> | Indian Ocean Deep Water   |
| <b>NPDW</b> | North Pacific Deep Water             | <b>BIW</b>  | Banda Intermediate Water  |
| <b>ACCS</b> | Antarctic Circumpolar Current System | <b>NIIW</b> | Northwest Indian Intermediate Water                             |
| <b>CDW</b>  | Circumpolar Deep Water               |             |   |



GFDL CM 2.4



Sea Surface Temperature ( $^{\circ}\text{C}$ )

